



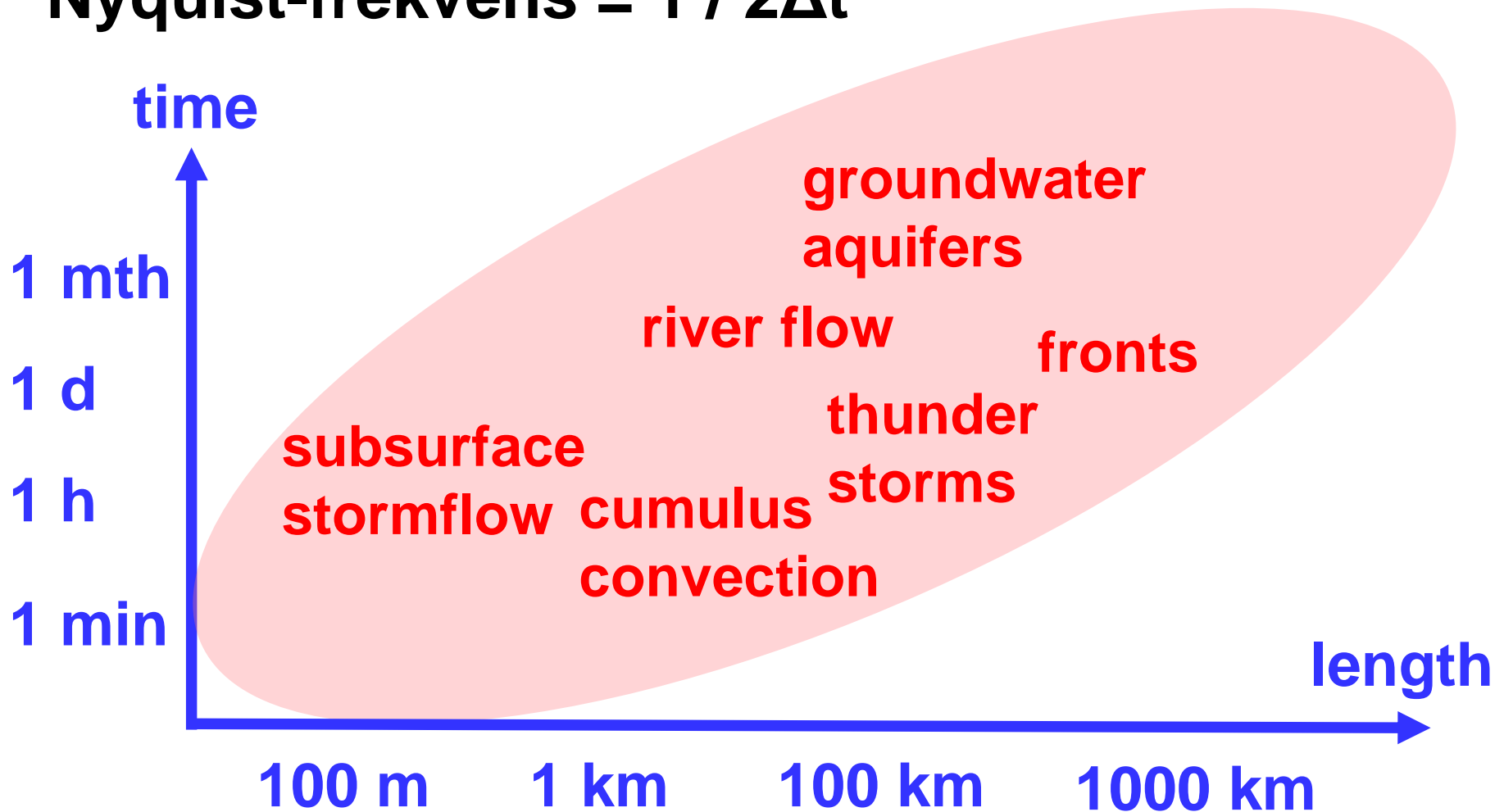
# Modellering av hydrologiske prosesser med høy oppløsning i tid og rom

Stein Beldring  
Norges vassdrags- og energidirektorat

# Hvorfor høy oppløsning i tid og rom?

- Modellens skala må gi en realistisk beskrivelse av prosessenes variasjon i tid og rom
- I nedbørfelt mindre enn  $\sim 10^2$  km<sup>2</sup> er reaksjonstiden i vassdraget med elver og innsjøer og det omkringliggende terrenget med vegetasjon, løsmasser og berggrunn mindre enn 1 døgn
- Den hydrologiske modellen må benytte tidsskritt som er kortere enn 1 døgn, for eksempel 1 time
- De meteorologiske dataene som driver den hydrologiske modellen må ha samme tidsoppløsning

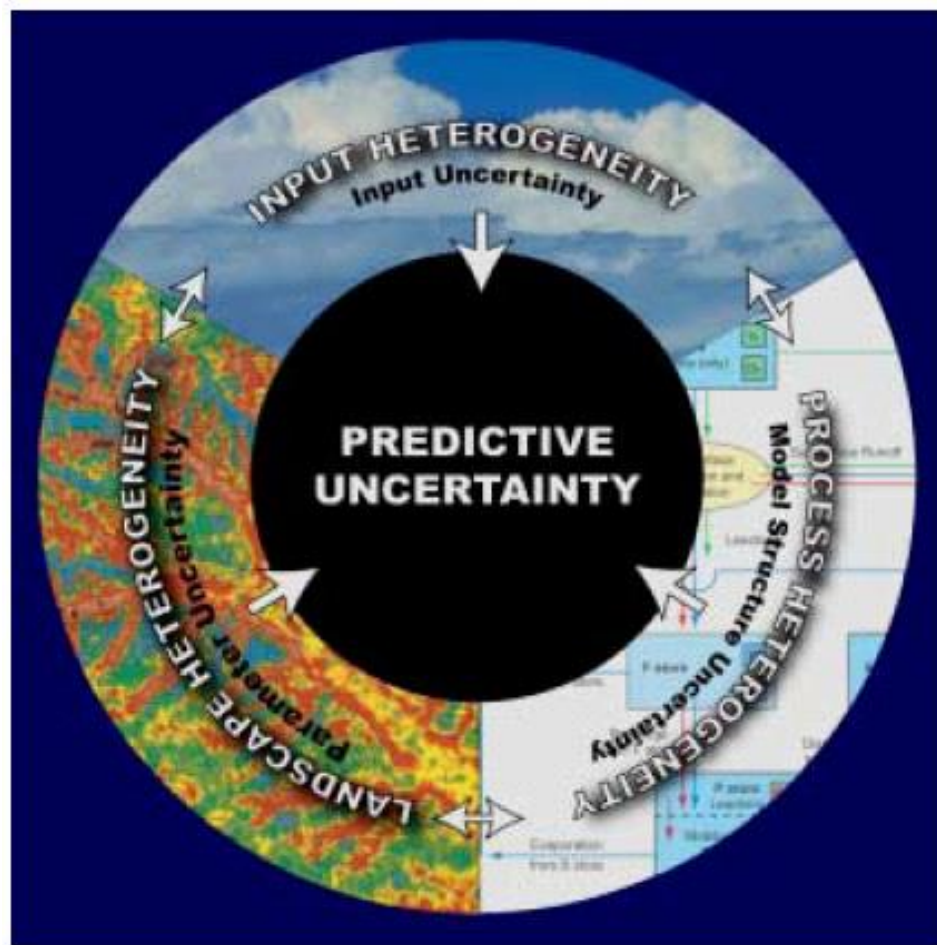
**Skala – en karakteristisk lengde eller tid for en prosess: (1) utstrekning eller varighet, (2) periode, (3) korrelasjonslengde eller -tid**  
**Nyquist-frekvens =  $1 / 2\Delta t$**



# Hydrologiske modeller - feilkilder

Atmosfære

Landoverflate



Hydrologisk modell

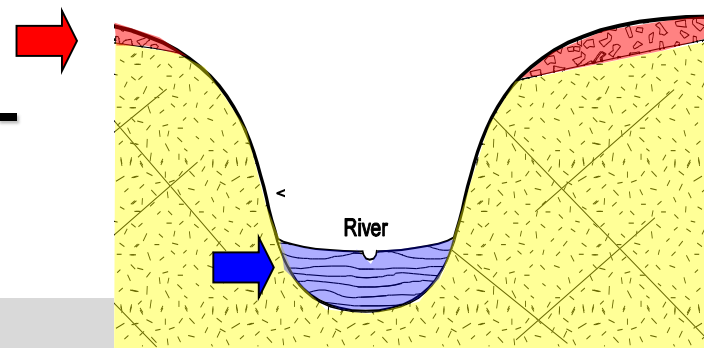
- parametrisering
- skala

Hydrologiske data



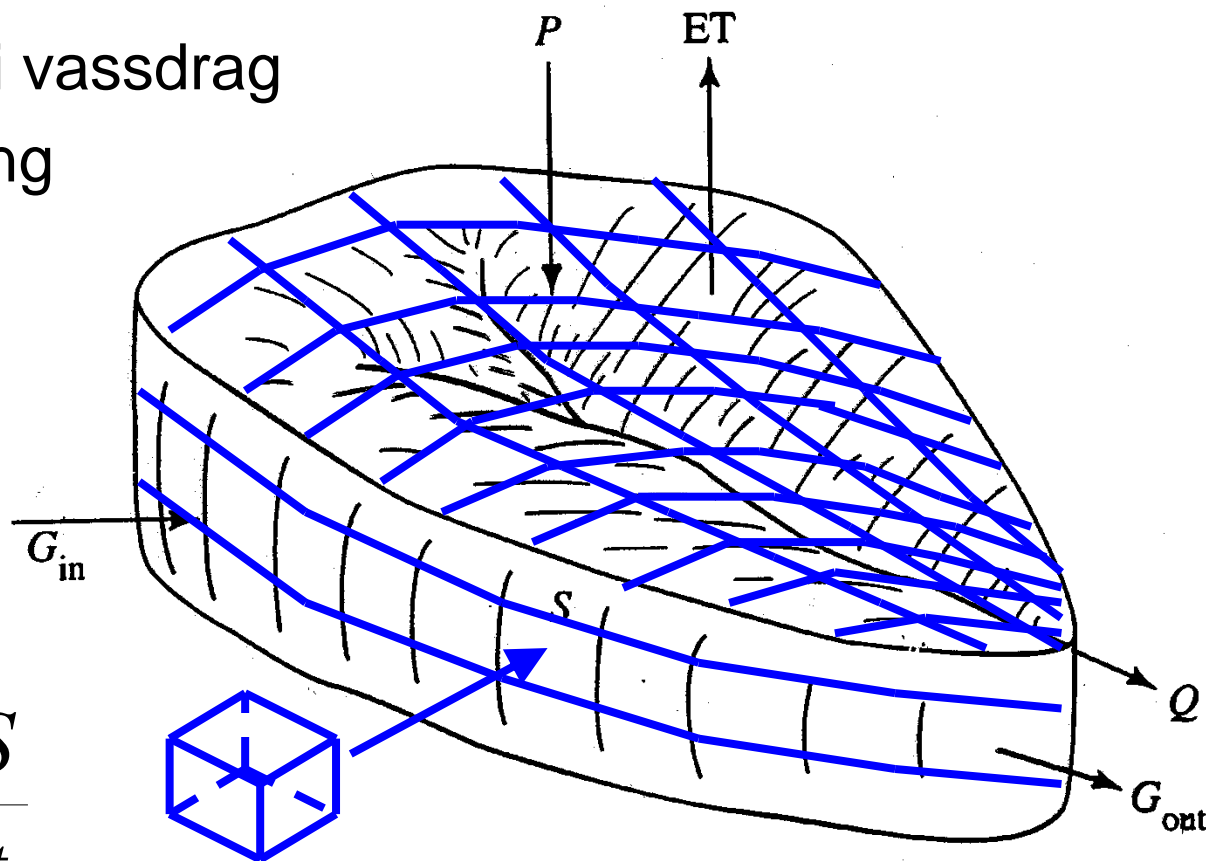
# Grunnvannsmagasinet i Norge består i hovedsak av:

- Små områder med fluviale og glasifluviale avsetninger med høy permeabilitet langs de største vassdragene og i lavlandet
- **Bunnmorene i skog- og fjellområder hvor grunnvann kontrolleres av nedbør**
- Oppsprukket berggrunn bestående av magmatiske og metamorfe bergarter uten primær porøsitet
- Noen få store, regionale grunnvannsmagasiner med høy porøsitet



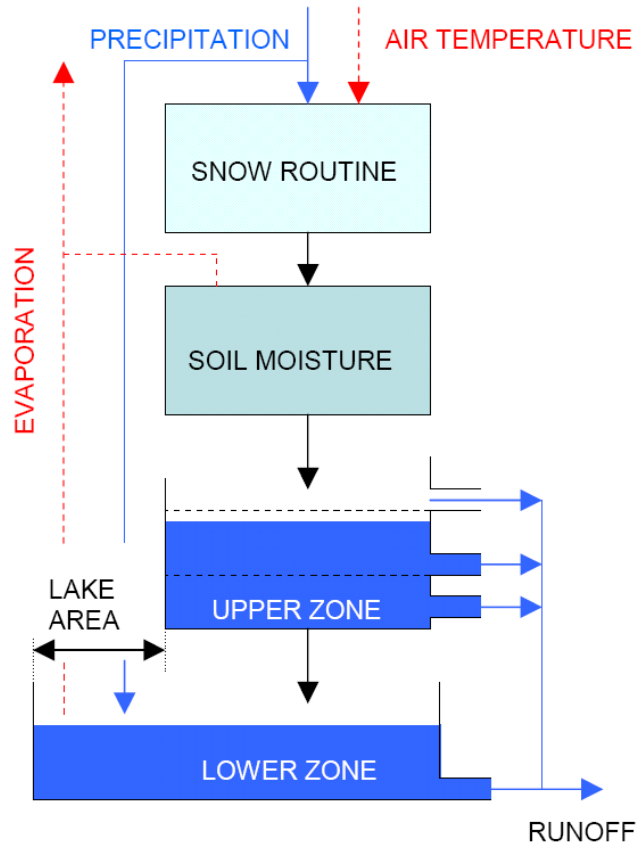
- P nedbør
- ET fordampning og transpirasjon
- G grunnvann
- Q vannføring i vassdrag
- S magasinerings

$$P + G_{in} = Q + G_{out} + ET + \Delta S$$



$$i - q = \frac{dS}{dt}$$

# The HBV-model – A Lumped Precipitation-Runoff model



**HBV** is an acronym formed from **H**ydrologiske **B**yrån avdeling for **V**attenbalans at SMHI, Sweden

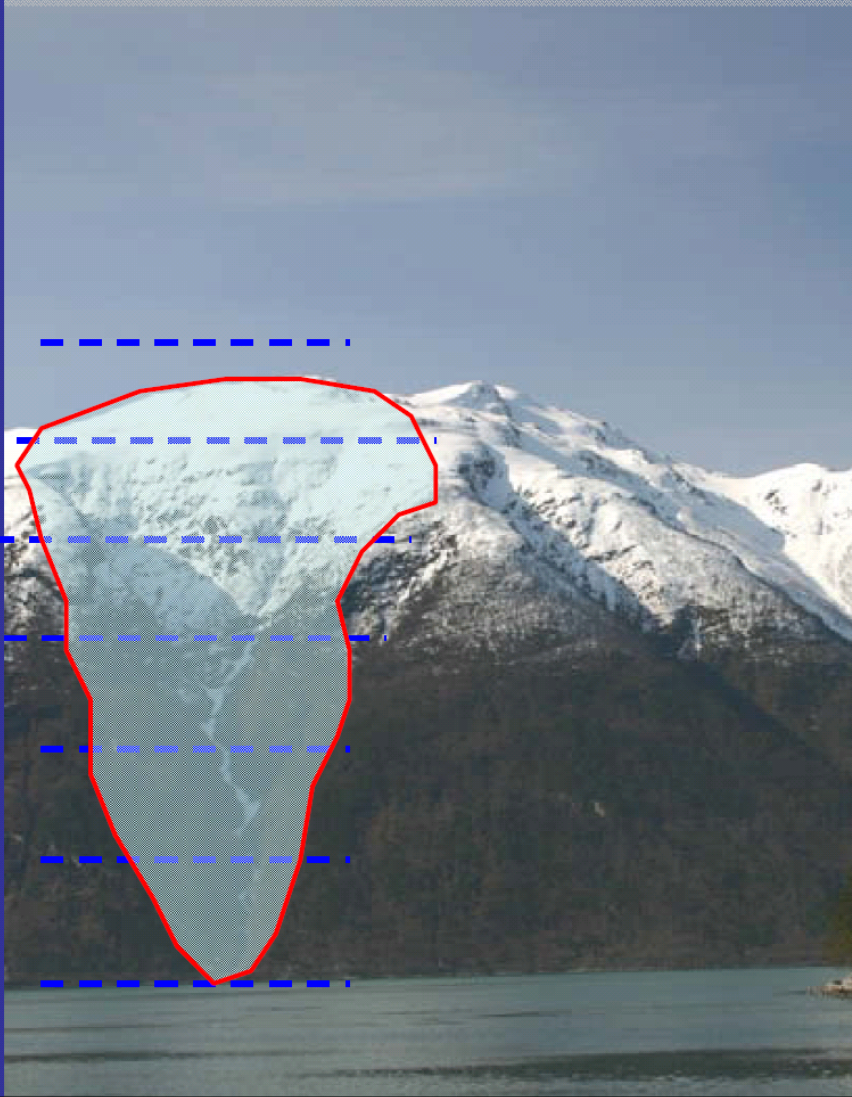
The model structure was made by **Sten Bergström**,

**Swedish Meteorological  
and Hydrological Institute**

The model is extensively used for making runoff/inflow forecasts to hydropower systems in Norway, Sweden and Finland, and many other countries in Europe (and outside)



# Snow routine must consider the effect of topography



More snow accumulates at higher elevation due to higher precipitation and lower air temperature.

Precipitation type and amount is a function of elevation

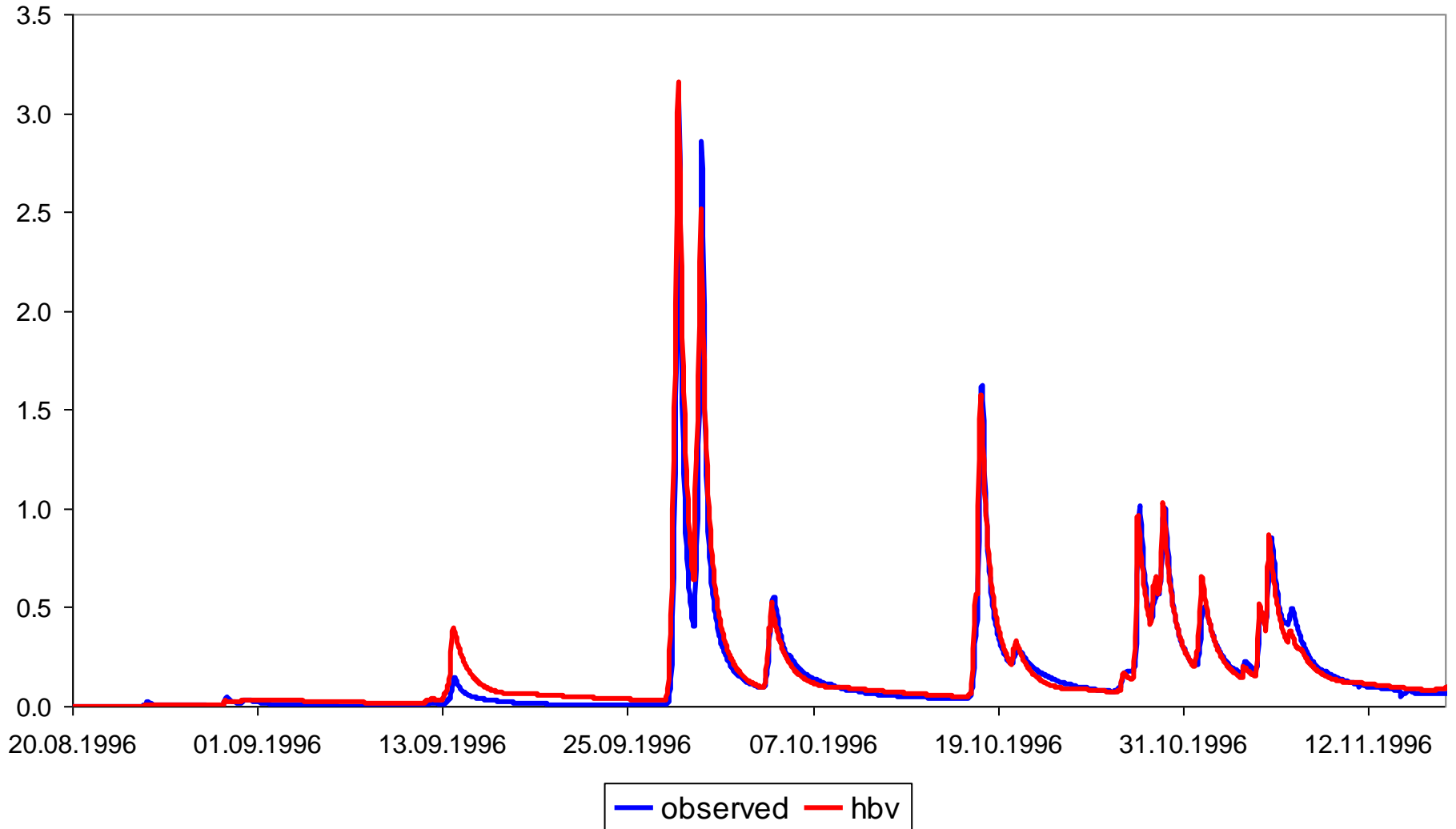
Also snow melt varies with elevation

Therefore, we must divide the catchments into a number of zones (typically 5-10 zones) from the lowest to highest levels, and compute accumulation and melt of snow in each zone



# HBV-modell med 10 høydesoner og tidsoppløsning 1 time

Discharge from Sæternbekken catchment ( $\text{m}^3/\text{s}$ )



# Kinematic Wave model *Beldring, 2002*

## Saturated subsurface flow

Hydraulic conductivity

$$K(u) = K_0 e^{au}$$

Darcy's law

$$q(u) = K(u) \sin \alpha$$

$$q(s) = \frac{K_0}{a} \sin \alpha e^{aD} (1 - e^{-as})$$

Continuity equation

$$\frac{\partial q}{\partial l} + \varepsilon \frac{\partial s}{\partial t} = i$$

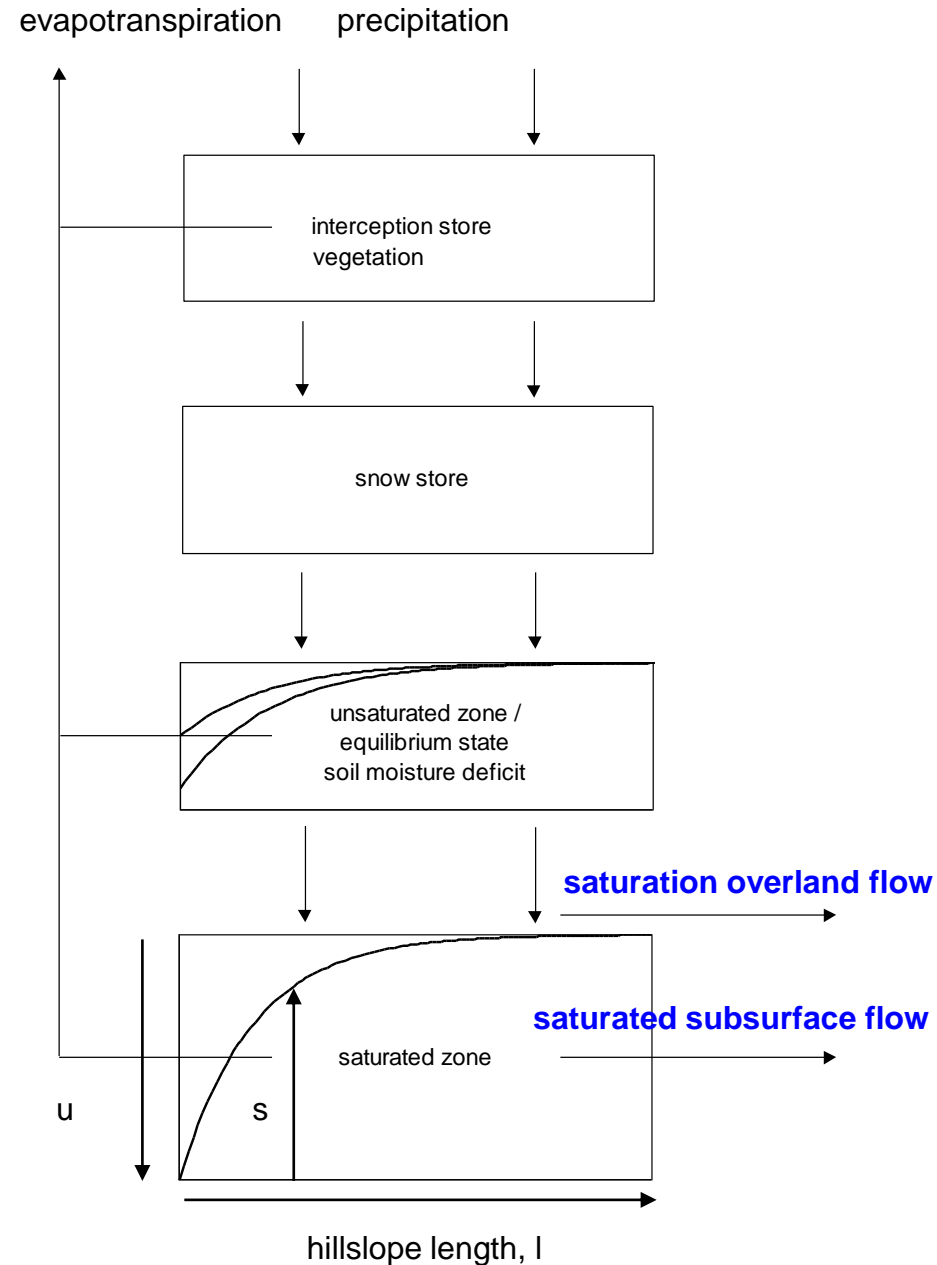
## Saturation overland flow

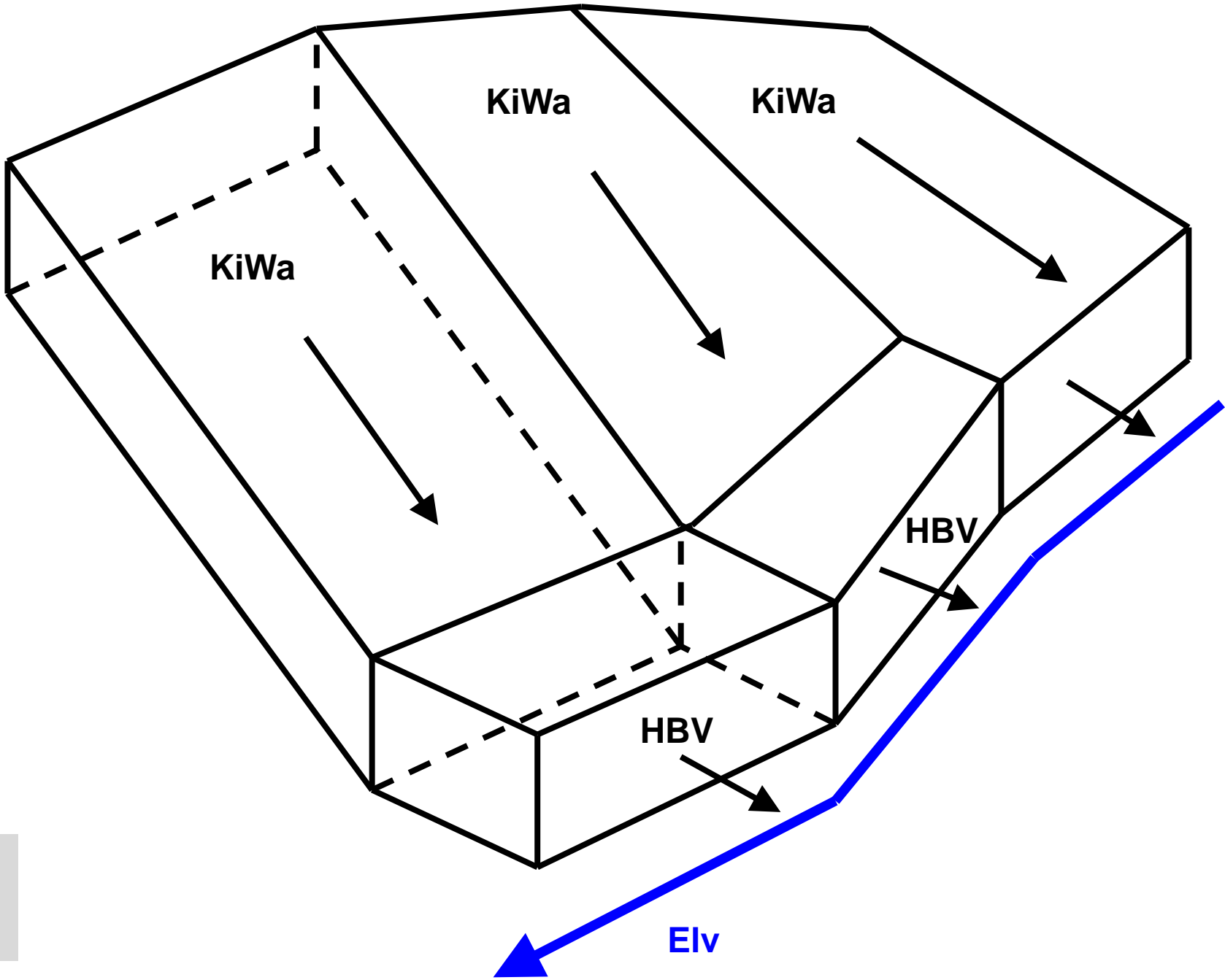
Overland flow flux law

$$p = \beta y^m$$

Continuity equation

$$\frac{\partial p}{\partial x} + \frac{\partial y}{\partial t} = i$$





# Romlig fordelt HBV-modell med 1 km<sup>2</sup> grid-ruter:

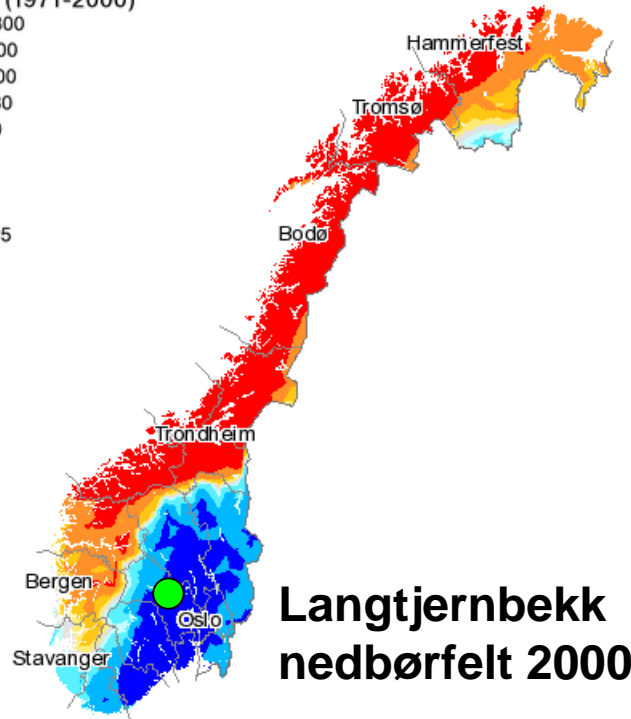
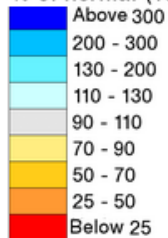
Månedlig nedbørsum og daglig grunnvannsmagasin høsten 2000 i forhold til gjennomsnittlig verdi for samme måned og dato

<http://www.xgeo.no/>

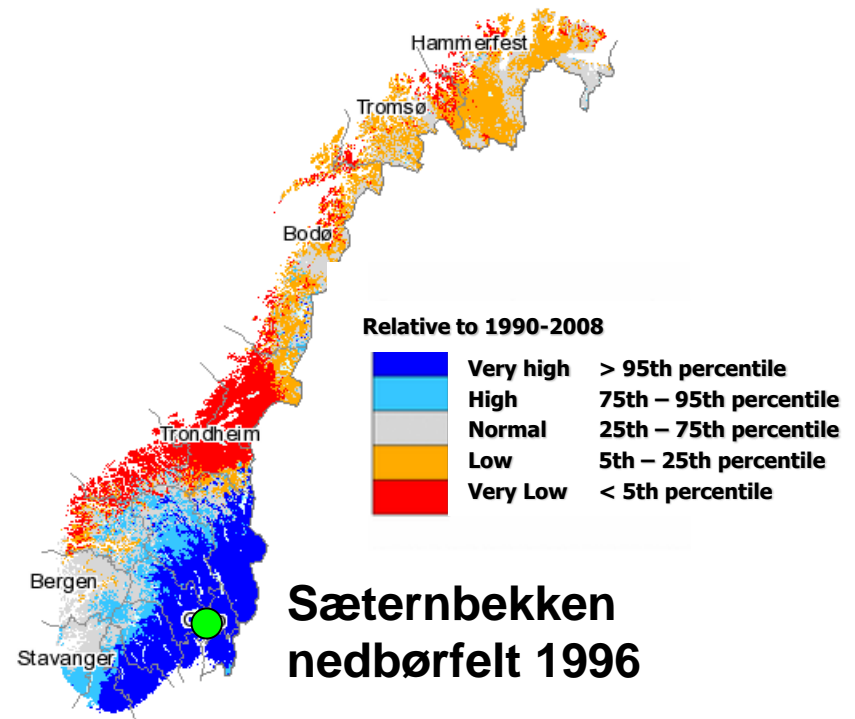
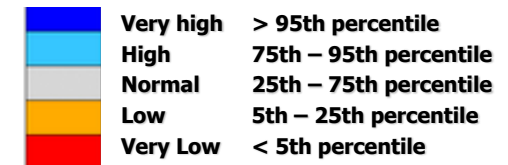
Precipitation monthly deviation (11.2000)

Groundwater level compared to normal (15.11.2000)

% of normal (1971-2000)



Relative to 1990-2008





# Sæternbekken

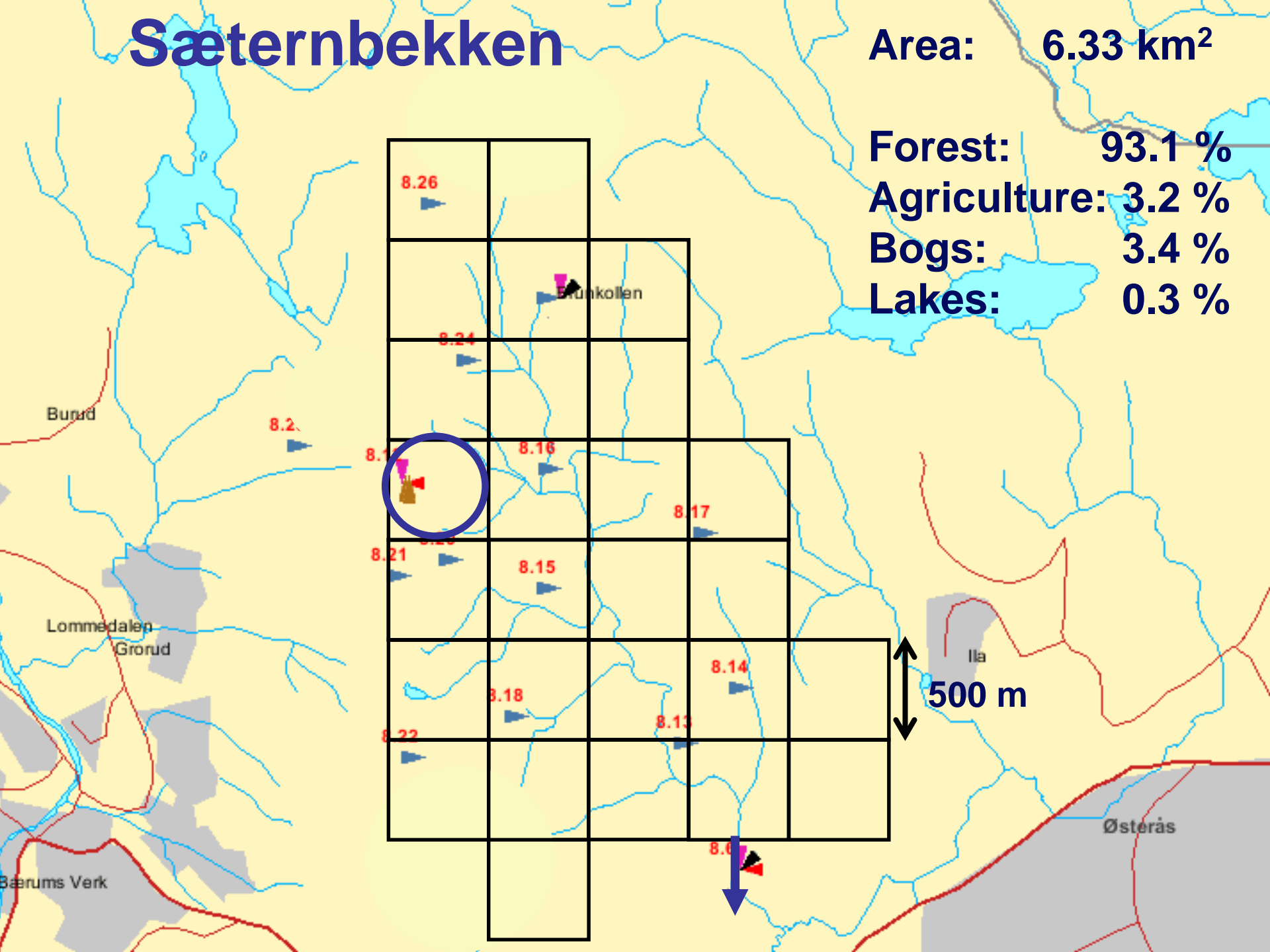
Area: 6.33 km<sup>2</sup>

Forest: 93.1 %

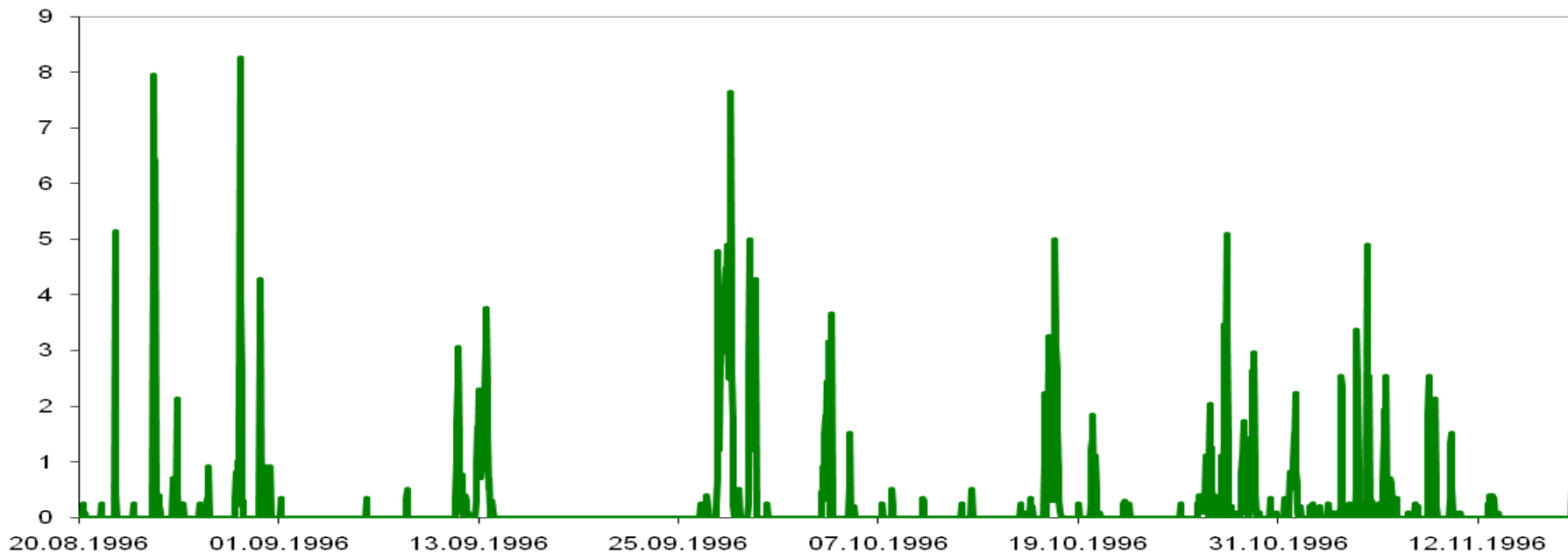
Agriculture: 3.2 %

Bogs: 3.4 %

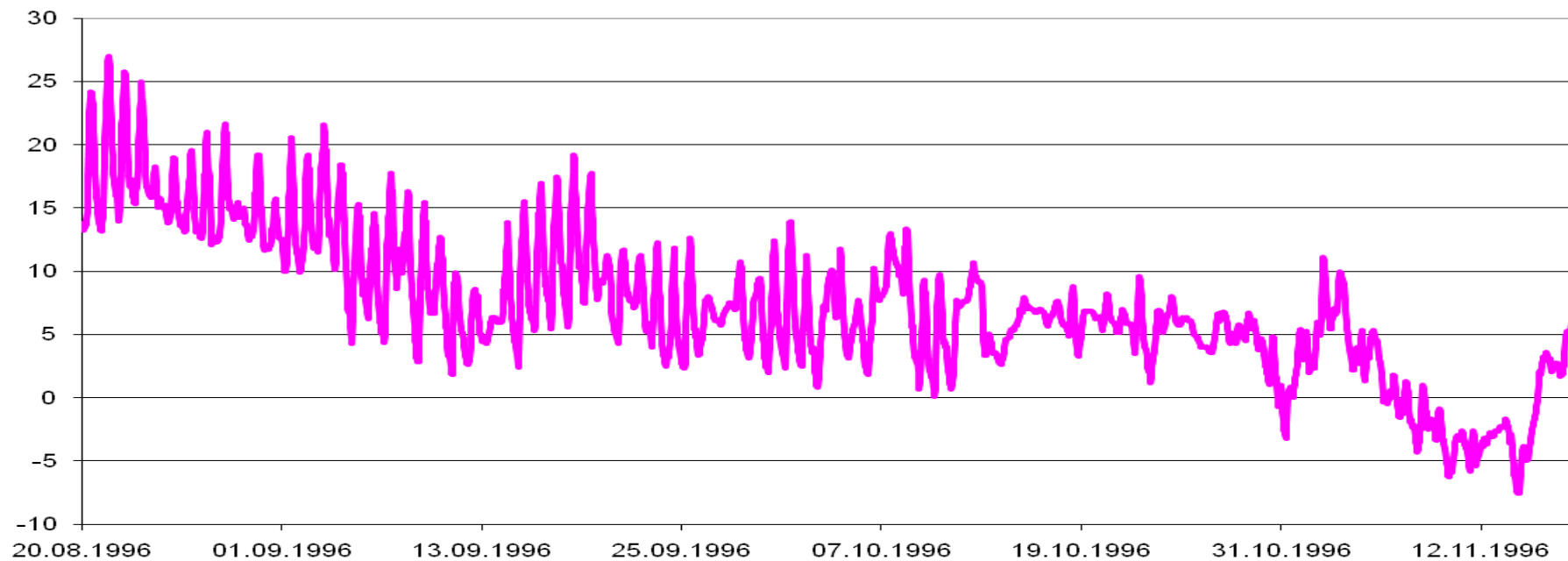
Lakes: 0.3 %



**Precipitation for Sæternbekken catchment (mm)**

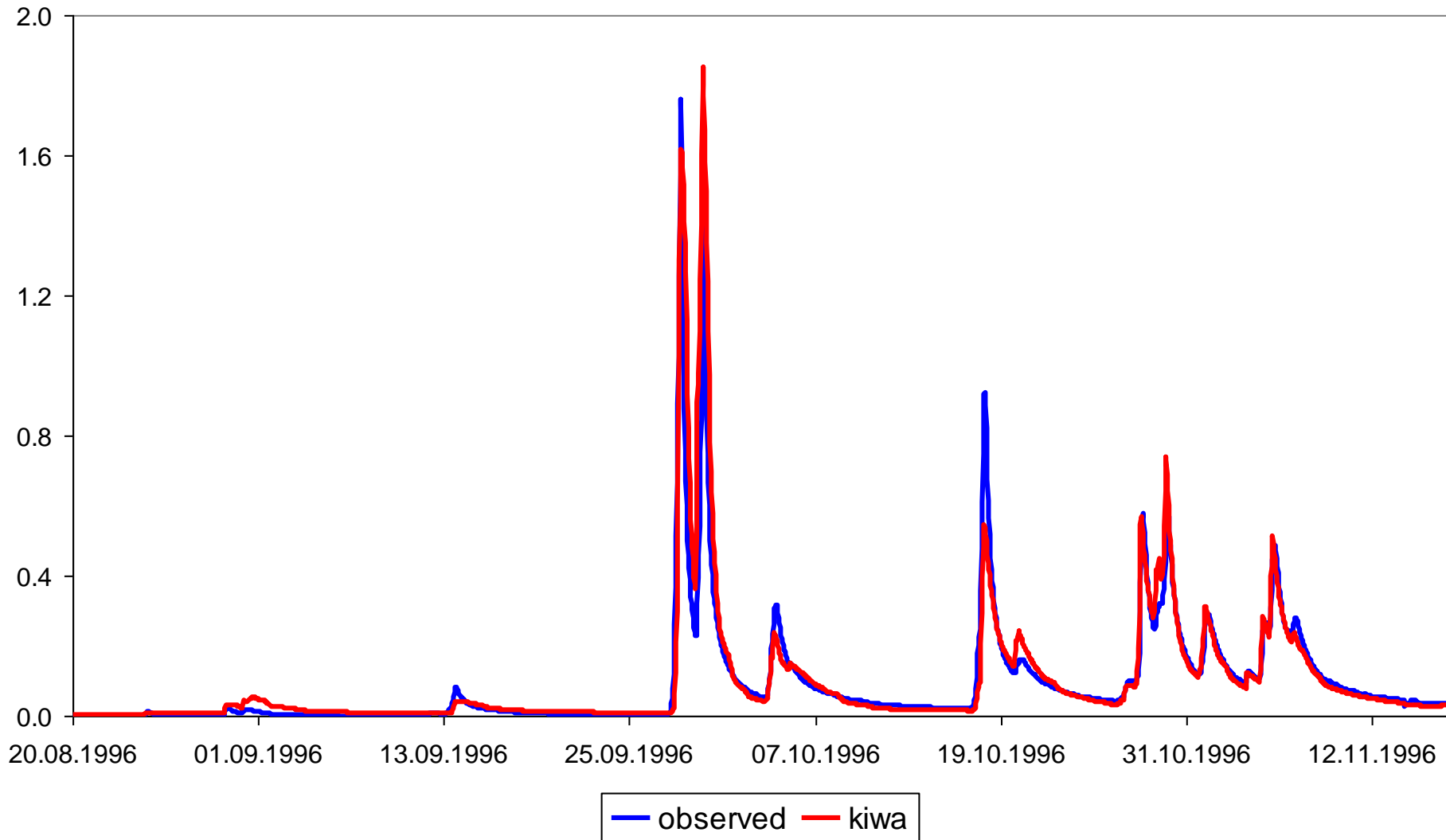


**Temperature for Sæternbekken catchment (°C)**

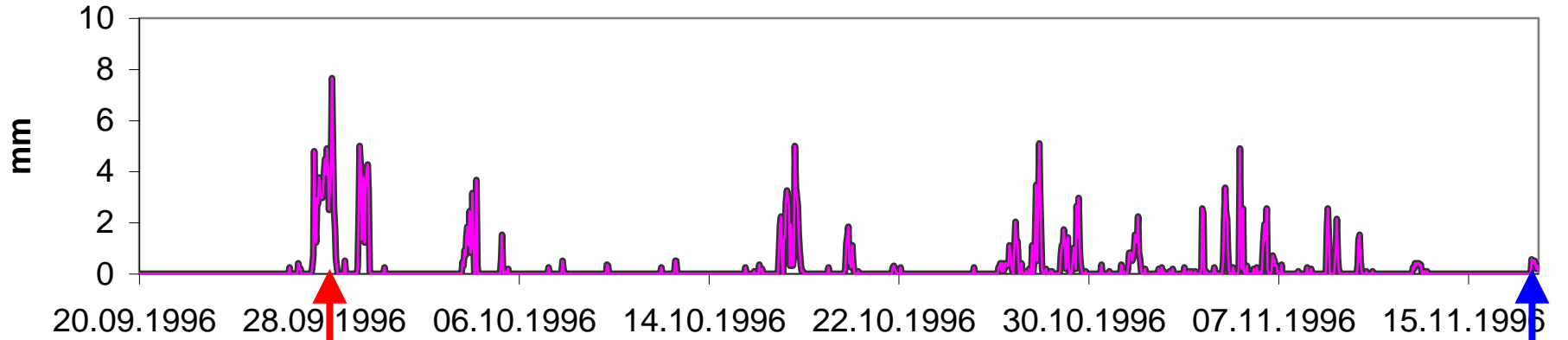


# Romlig fordelt kinematic-wave-algoritme med skråninger (hillslopes) på ca. 0.25 km<sup>2</sup> som beregnings-element, routing i elvenett og tidsoppløsning 1 time

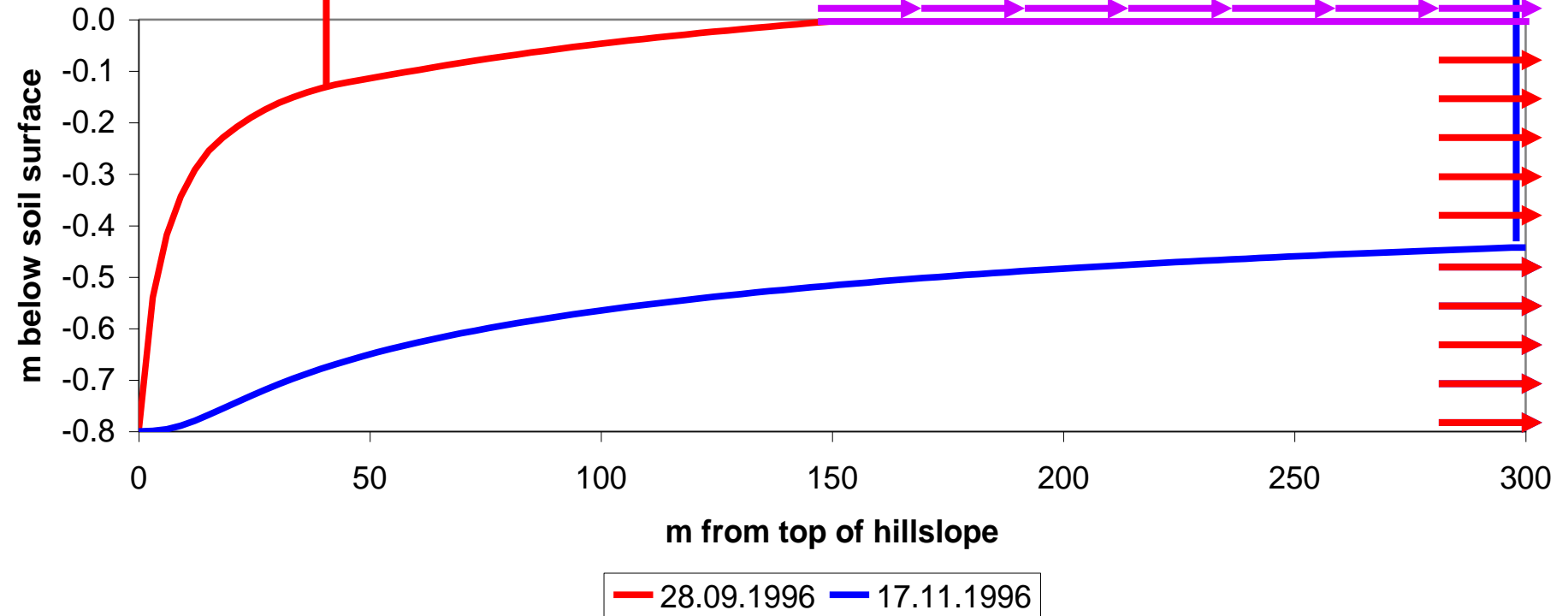
Runoff from Sæternbekken catchment (mm)



# Observed precipitation in Sæternbekken catchment

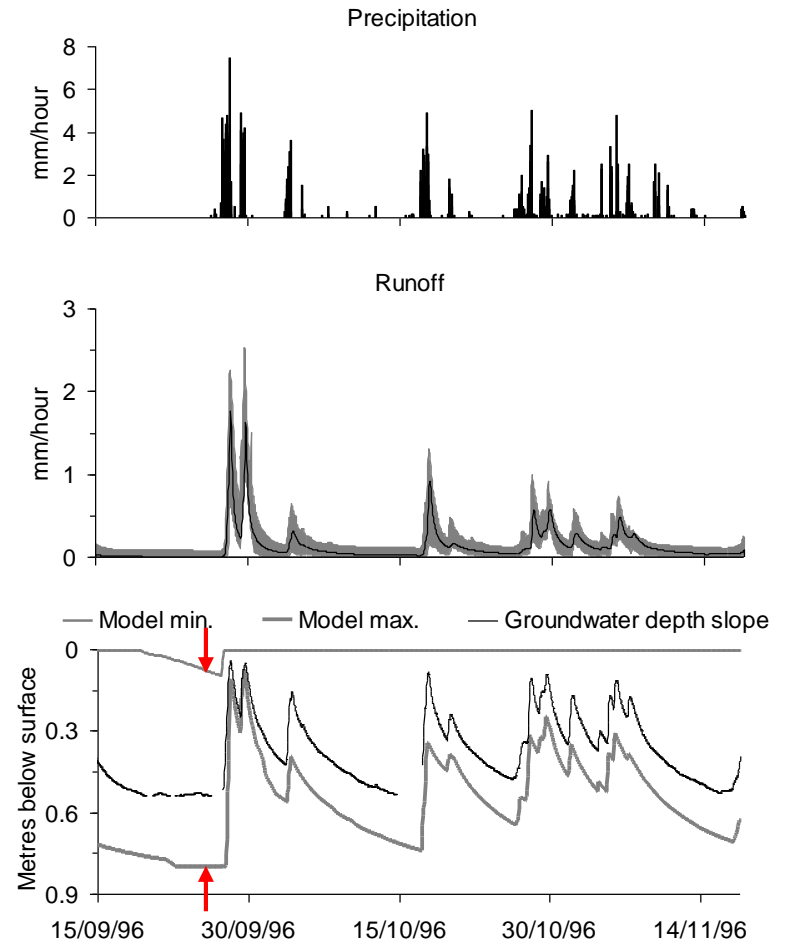
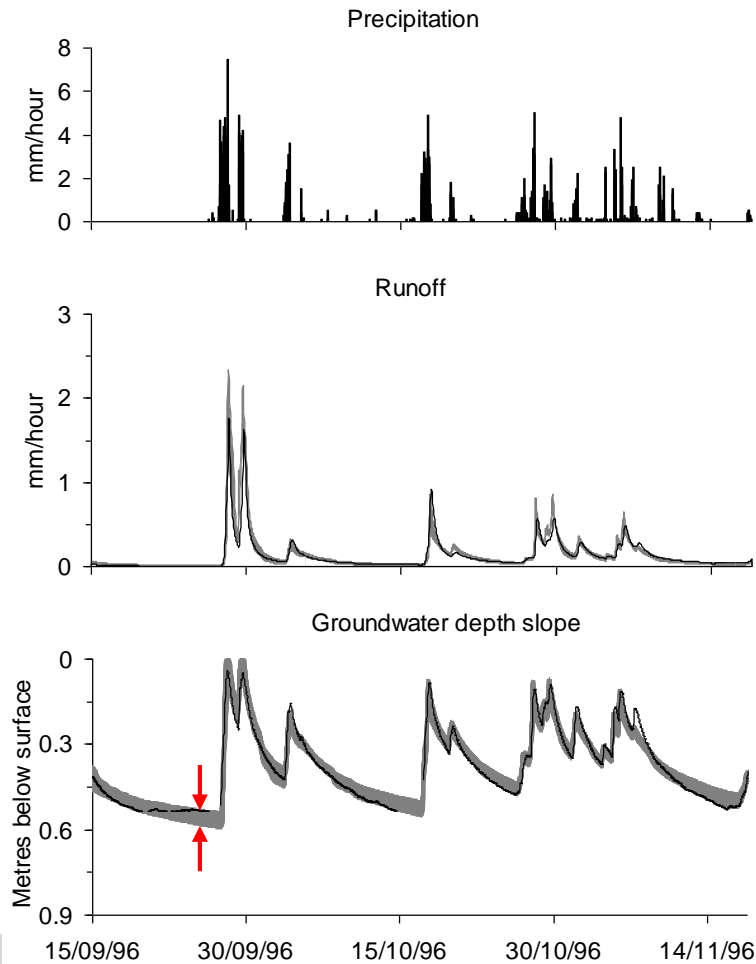


## KinematicWave model for Sæternbekken catchment: Simulated groundwater table depth vs. length from top of hillslope





# Kinematic Wave model with hourly time step for Sæternbekken catchment



**Multi-objective calibration**  
**Runoff and groundwater**

**Single-objective calibration**  
**Runoff**

## Langtjernbekk catchment Landscape elements and stream network

**Area:** 4.81 km<sup>2</sup>

**Forest:** 87.3 %

**Bogs:** 5.8 %

**Lakes:** 6.9 %

**Hillslopes:** Kinematic Wave

**Bogs:** HBV

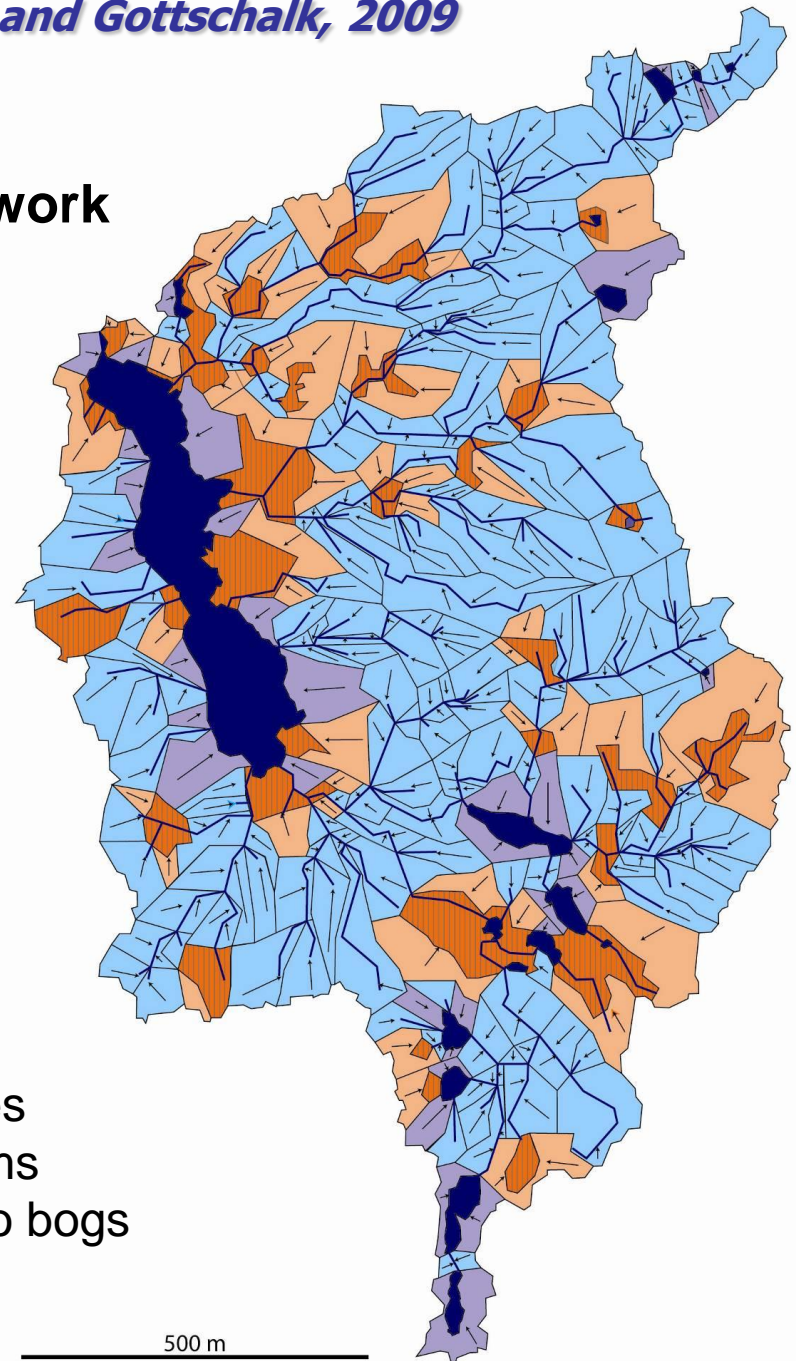
**Dark blue** : lakes

**Purple** : KiWa discharging into lakes

**Blue** : KiWa discharging into streams

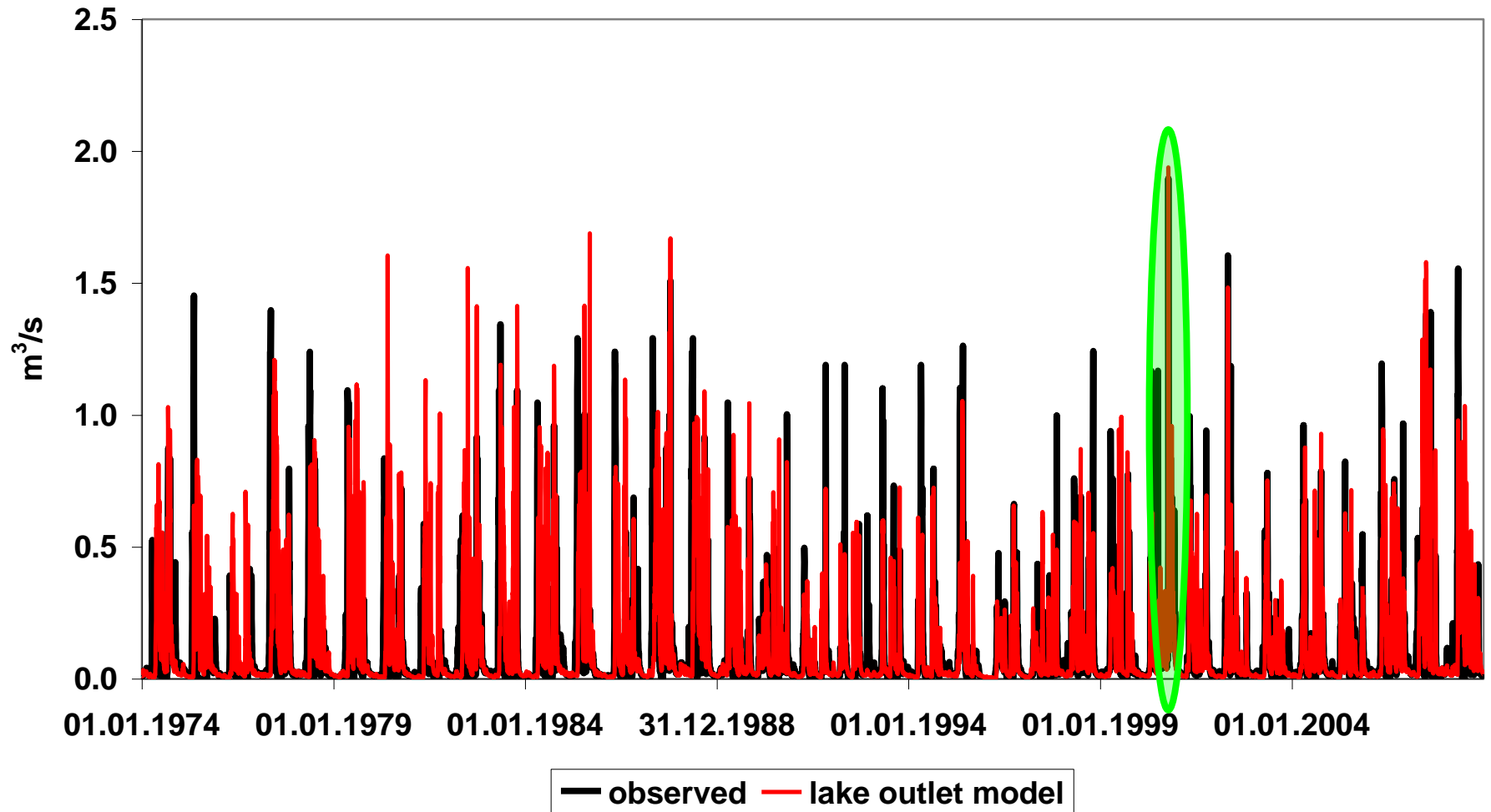
**Light brown**: KiWa discharging into bogs

**Dark brown** : HBV - bogs



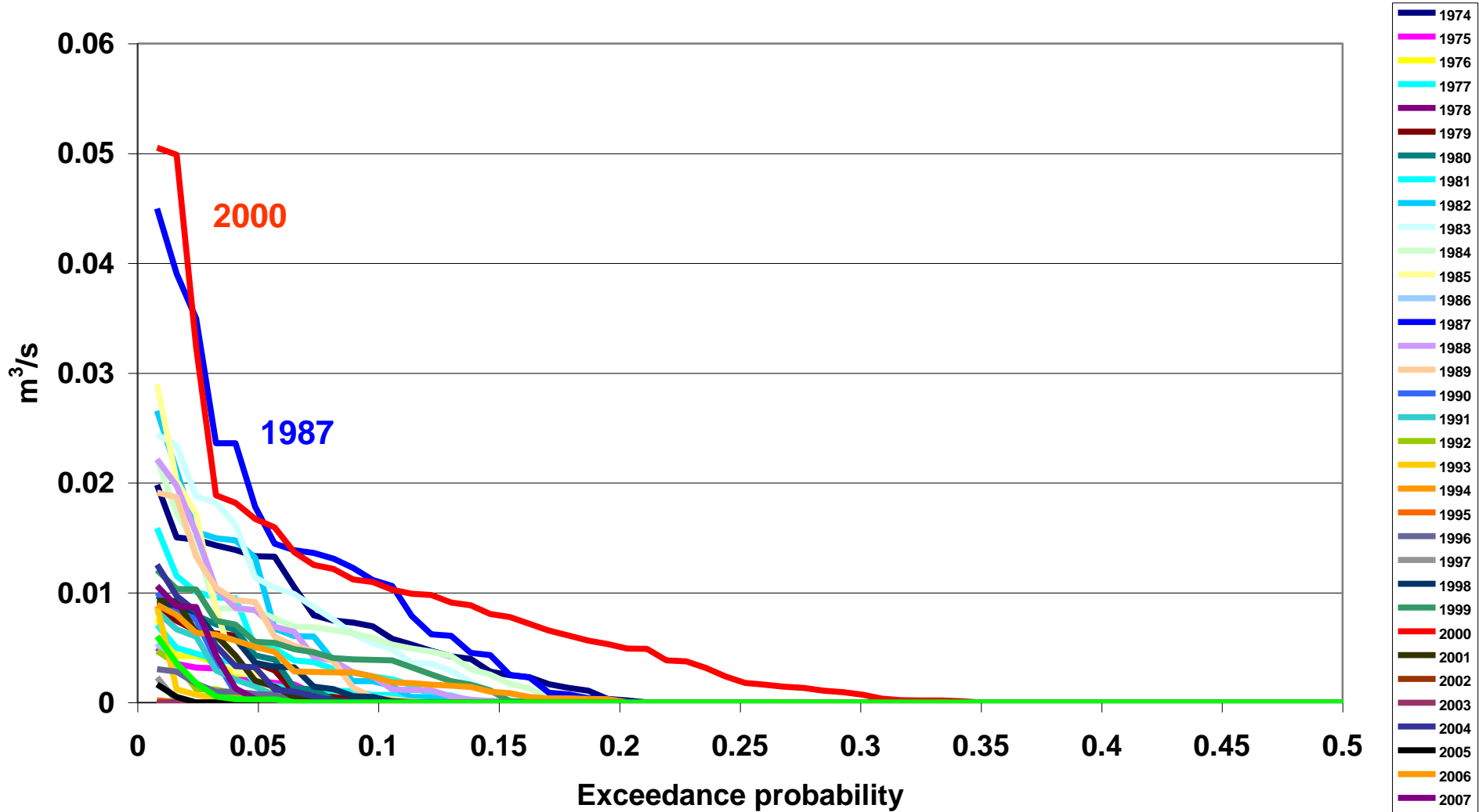
# KiWa model for Langtjernbekk catchment

Discharge from Langtjernbekk catchment  
time step: 1 day



# KiWa model for Langtjernbekk catchment

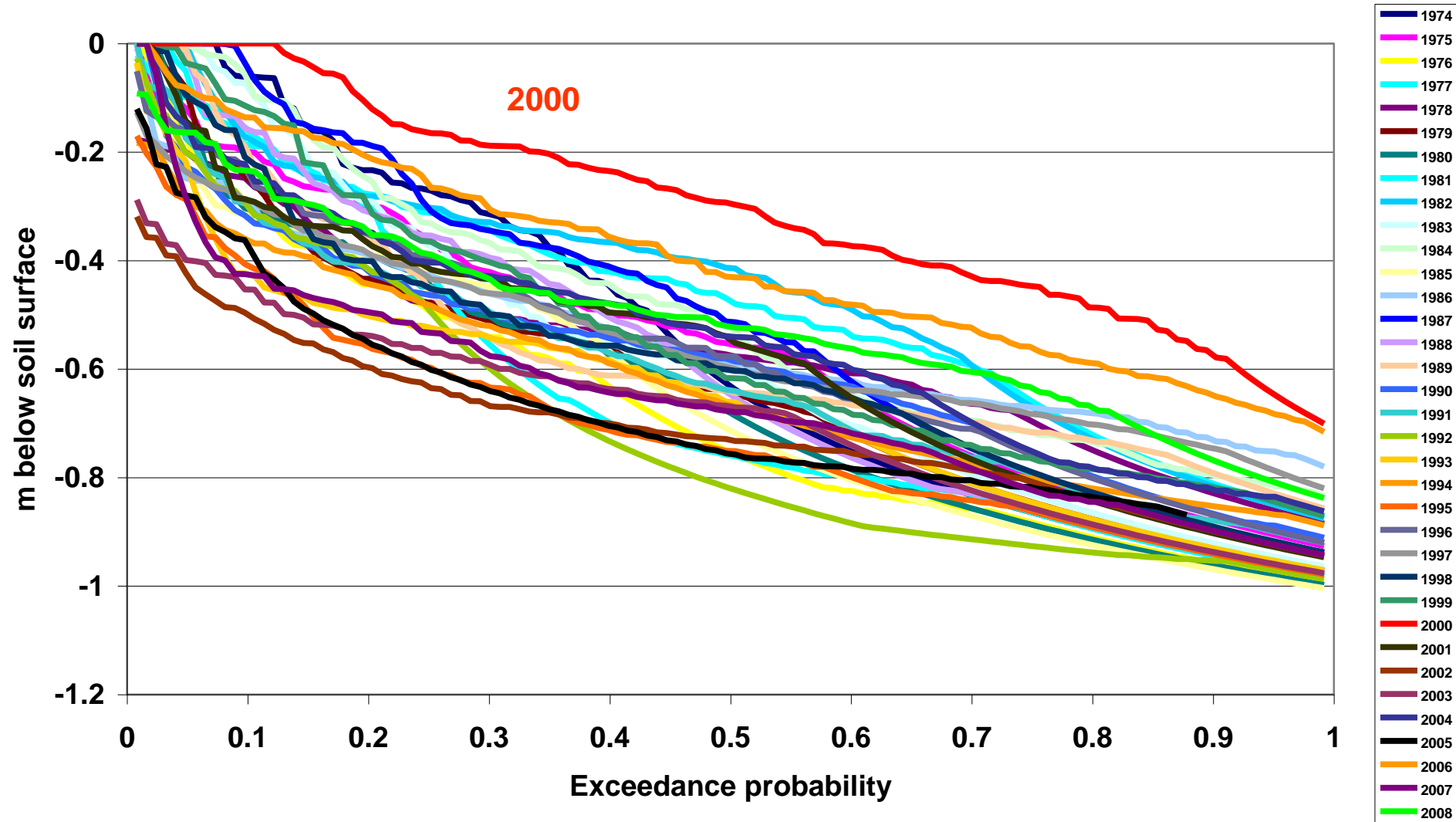
Cumulative distribution function of daily overland flow for Sep. - Dec.





# KiWa model algorithm for Langtjernbekk catchment

Cumulative distribution function of groundwater table depth for Sep. - Dec.



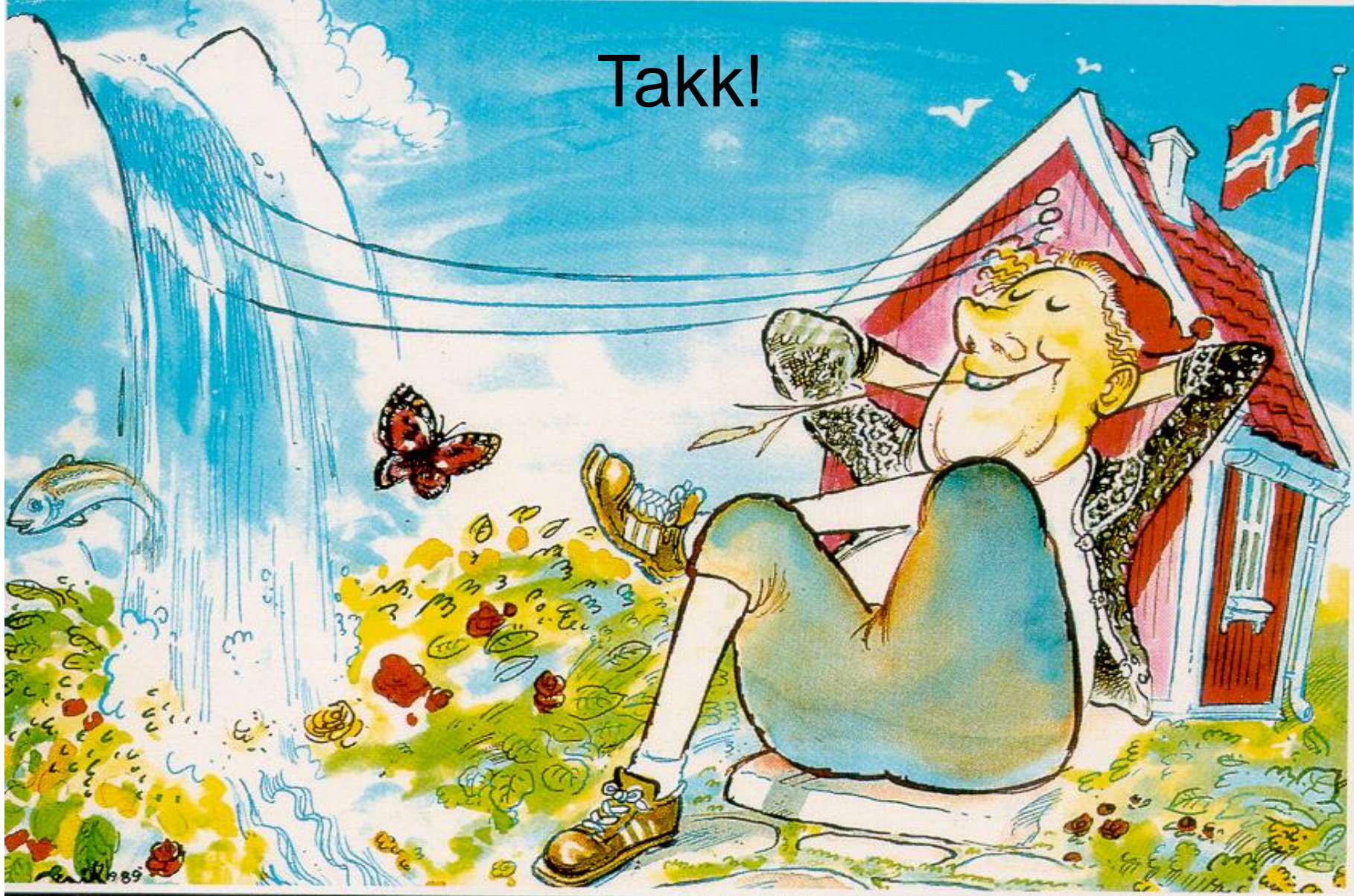
# Skala i rom og tid er viktig ved valg av modellstruktur og data

- Vannføring fra små nedbørfelt blir bedre beskrevet med hydrologiske modeller med fin tidsoppløsning
- Romlig diskretisering av modellens beregningselement har betydning for detaljeringsgrad i beskrivelse av prosesser
- Algoritmene som benyttes i den hydrologiske modellen må velges etter formål og behov – modellen må virke godt av de riktige grunnene
- Modellens oppløsning i tid og rom må ta hensyn til skala i de dominerende prosessene
- Prosess-skala, data-skala og modell-skala må være i overensstemmelse

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Takk!



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