

Norwegian University  
of Life Sciences



Arungen Lake, As, Norway.



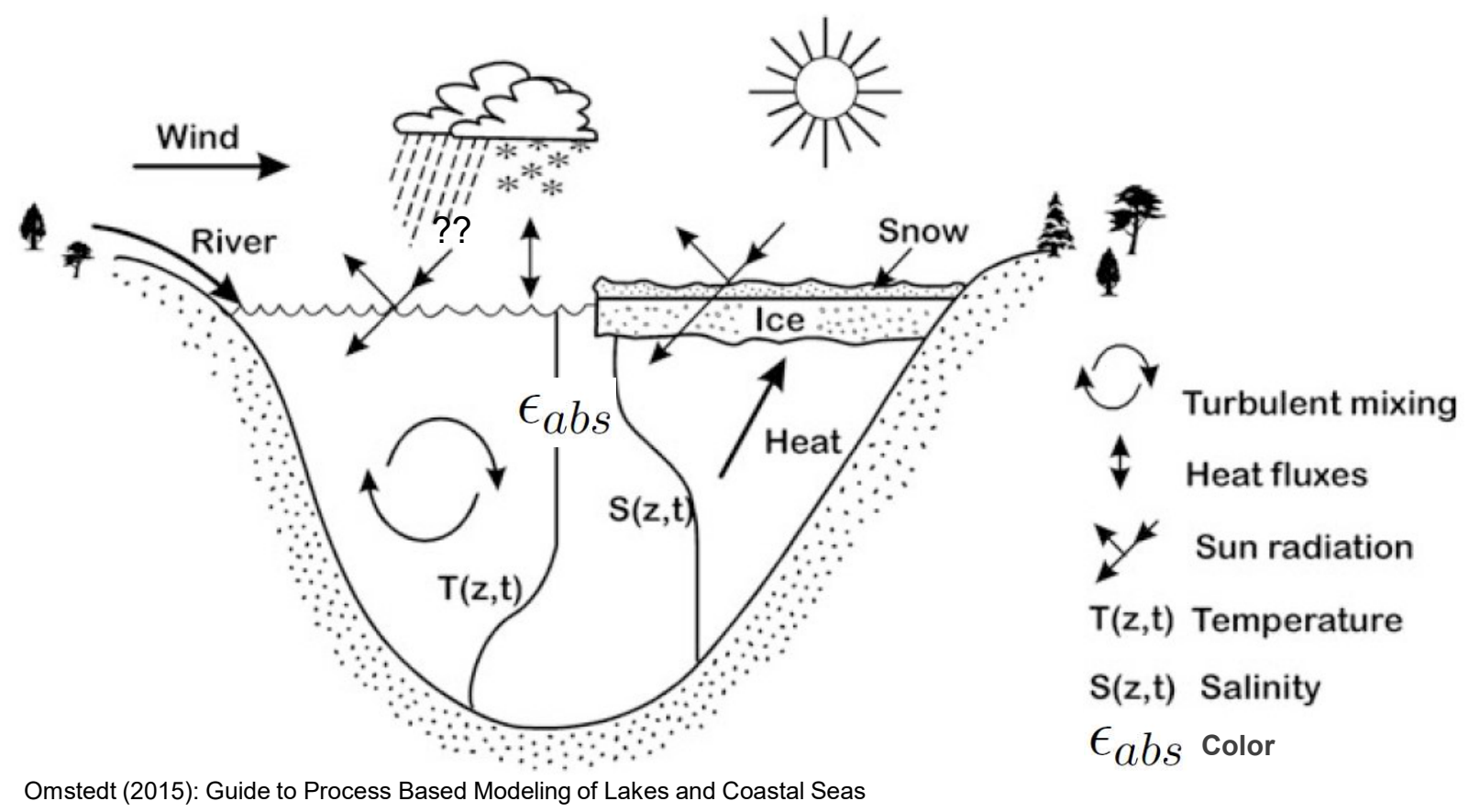
# Lake temperature and global warming

Sverre Anmarkrud<sup>1</sup>, Alemayehu (Alex) Adugna Arara<sup>2</sup>, Anne Kværnø<sup>3</sup>, and Nils-Otto Kitterød<sup>1</sup>

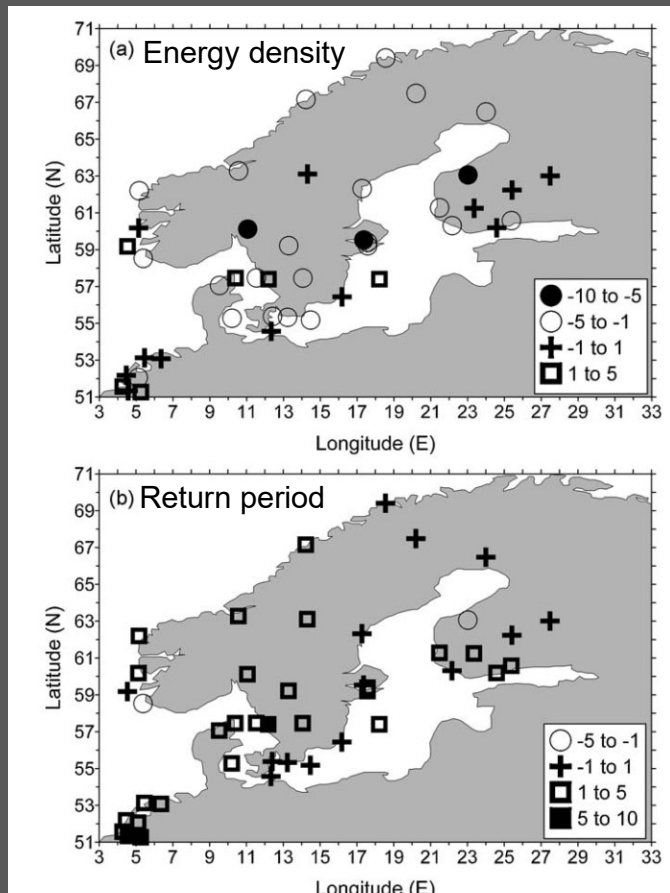
- 1) Faculty of Environmental Sciences and Natural Resource Management, Norwegian University of Life Sciences, Ås, Norway.
- 2) Department of Mathematics, Hawassa University, Ethiopia.
- 3) Department of Mathematics, Norwegian University of Science and Technology, Trondheim, Norway.

The Norwegian Hydrological Council/Norsk hydrologiråd  
5<sup>th</sup> Conference on Modelling Hydrology Climate and Land Surface Processes  
September 17 – 19, 2019, Lillehammer, Norway

# Lake temperature



# Wind velocities and global warming?



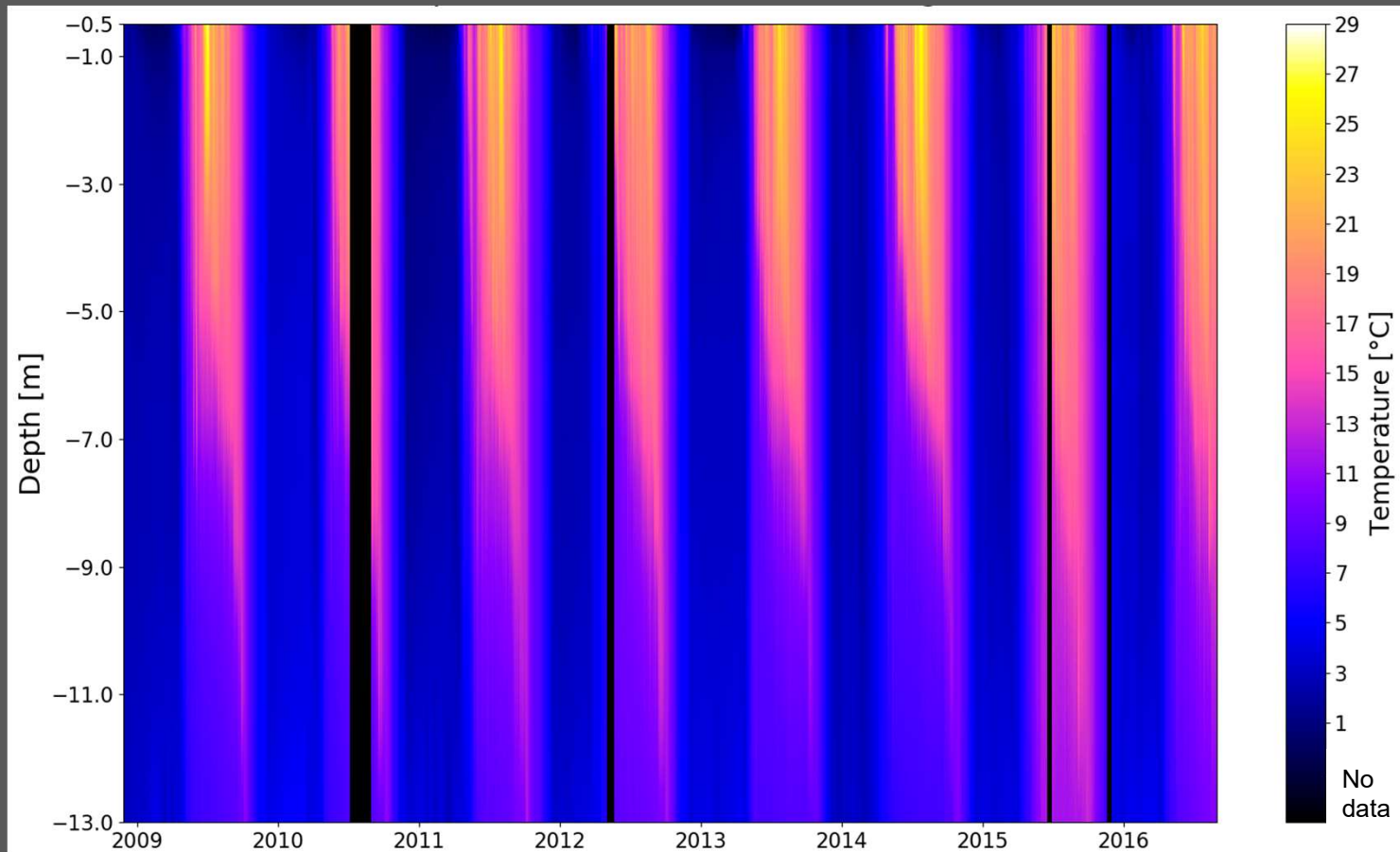
$$E = \frac{1}{2} \rho U^3$$

$E$  = energy density ( $\text{W m}^{-2}$ ),  
 $\rho$  = air density ( $\text{kg m}^{-3}$ ),  
 $U$  = wind speed at hub-height ( $\text{m s}^{-1}$ ).

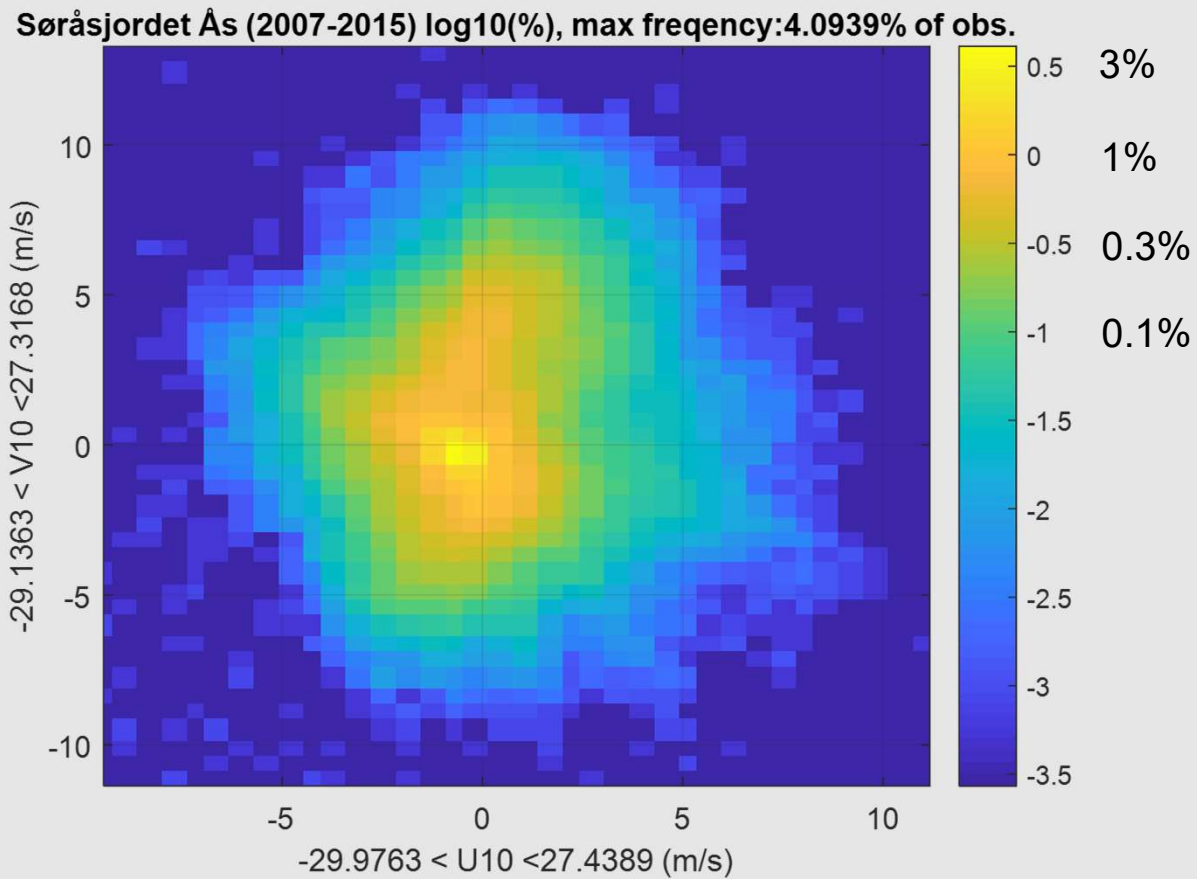
Ensemble average difference in percent of (a) energy density and (b) 50-year return period wind speed computed using a probabilistic empirical downscaling Approach for 43 stations across northern Europe based on output from 8-GCMs (BCCR-BCM2.0, CGCM3.1, CNRM-CM3, ECHAM5/MPI-OM, GFDL-CM2.0, GISS-ModelE20/Russell, IPSL-CM4, and MRI-CGCM2.3.2.). The future time period is 2081–2100, while the historical period is 1961–1990. **A positive value indicates higher energy density or extreme wind in the later time period**



# Temperature observations from lake Årungen



Probability density function of wind velocities from Søråsjordet in the period from 2007 to 2015, 10 min average.





# The Simstrat Model

<https://www.eawag.ch/en/departement/surf/projects/simstrat/>

- 1D physical lake model.
- For simulation of stratification and mixing in deep lakes.
- Originally developed by Goudsmit et al. (2002).
- Applied to lakes with different physical properties.
- Open source: The Simstrat code is freely available on Github:
  - <https://github.com/Eawag-AppliedSystemAnalysis/Simstrat/>



# The Main properties of the Simstrat Model

<https://www.eawag.ch/en/departement/surf/projects/simstrat/>



- Simulates temperature, salinity, stratification and ice cover.
- k- $\epsilon$  model for turbulent mixing and buoyancy oscillations.
- Includes energy transfer to mixing via internal seiches
- Hydrology: inflows can be added at specified depths or with density-dependent intrusions.
- Variable lake surface levels.
- Programmed in object-oriented Fortran 2003.
- Parameter estimation using PEST is implemented

## Basic Model Equations

The basic set of equations for temperature  $T$ , mean horizontal velocity components  $u$  and  $v$  with respect to  $x$  and  $y$ , turbulent kinetic energy (TKE) per unit mass  $k$  and the TKE dissipation rate  $\epsilon$  are

$$\frac{\partial T}{\partial t} = \frac{1}{A} \frac{\partial}{\partial z} \left( A(\nu'_t + \nu') \frac{\partial T}{\partial z} \right) + \frac{1}{\rho_0 c_p} \frac{\partial H_{sol}}{\partial z} + \frac{dA}{dz} \frac{H_{geo}}{A \rho_0 c_p}, \quad \nu'_t = c'_\mu \frac{k^2}{\epsilon}, \quad (1)$$

$$\frac{\partial u}{\partial t} = \frac{1}{A} \frac{\partial}{\partial z} \left( A(\nu_t + \nu) \frac{\partial u}{\partial z} \right) + f v, \quad \nu_t = c_\mu \frac{k^2}{\epsilon}, \quad (2)$$

$$\frac{\partial v}{\partial t} = \frac{1}{A} \frac{\partial}{\partial z} \left( A(\nu_t + \nu) \frac{\partial v}{\partial z} \right) - f u, \quad (3)$$

$$\frac{\partial k}{\partial t} = \frac{1}{A} \frac{\partial}{\partial z} \left( A \nu_k \frac{\partial k}{\partial z} \right) + P + P_{seiche} + B - \epsilon, \quad \nu_k = \frac{c_\mu}{\sigma_k} \frac{k^2}{\epsilon}, \quad (4)$$

$$\frac{\partial \epsilon}{\partial t} = \frac{1}{A} \frac{\partial}{\partial z} \left( A \nu_\epsilon \frac{\partial \epsilon}{\partial z} \right) + \frac{\epsilon}{k} (c_{\epsilon 1} (P + P_{Seiche}) + c_{\epsilon 3} B - c_{\epsilon 2} \epsilon), \quad \nu_\epsilon = \frac{c_\mu}{\sigma_\epsilon} \frac{k^2}{\epsilon}, \quad (5)$$

$$P = \nu_t \left( \left( \frac{\partial u}{\partial z} \right)^2 + \left( \frac{\partial v}{\partial z} \right)^2 \right), \quad B = -\nu'_t N^2, \quad N^2 = -\frac{g}{\rho_0} \frac{\partial \rho}{\partial z}, \quad (6)$$

$$P_{Seiche}(z) = \frac{1 - 10\sqrt{C_{Defl}}}{\rho_0 c A_{boundary}} N^{2q} \frac{dA}{dz} \gamma E_{seiche}^{3/2}, \quad E(t) = \left( \frac{1}{E(t=0)} + \frac{\gamma}{2} t \right)^{-2}, \quad (7)$$



where

- $\rho_0$  is reference density,
- $c_p$  is specific heat of lake water,
- $A$  is the surface area of the lake at depth  $z$  and  $A_{boundary}$  is the total bottom area,
- $H_{sol} = H_{Sol,0}(1 - r_s)e^{-\epsilon_{abs}z}$  is the shortwave solar radiation penetrating the water,  $H_{Sol,0}$  is measured solar radiation above the water surface,  $r_s$  and  $\epsilon_{abs}$  are the reflection and extinction coefficient of the lake water
- $H_{geo}$  is the geothermal heat flux,
- $f$  is the Coriolis parameter,
- $\nu$  and  $\nu_t$  are molecular and turbulent viscosity (momentum),
- $\nu'_t$  and  $\nu'$  are molecular and turbulent diffusivity of temperature,
- $\nu_\epsilon$  and  $\nu_k$  are the turbulent diffusivities of energy dissipation and TKE,
- $P$  is the shear stress production,
- $B$  is the buoyancy production and  $N$  is the Brunt-Väisälä frequency,
- $P_{Seiche}$  is the production of TKE due to internal seiching,
- $c_{\epsilon 1}, c_{\epsilon 2}, c_{\epsilon 3}, c_\mu, c'_\mu, \sigma_k$  and  $\sigma_\epsilon$  are model constants,
- $\gamma$  is a constant of proportionality (can be estimated),
- $C_{Def}$  is effective bottom friction coefficient and
- $E_{Seiche}$  is total energy contained in the Seiche motion.



# The stochastic wind model

## 3.1 The model

Let  $Z(t) = [u(t), v(t)]$  be the vector representing the wind in two orthogonal directions (east-west and south-north). A stationary and two-points autocorrelated process can be defined by the following stochastic linear differential equations:

$$dZ = AZdt + BdW \quad (20)$$

where  $A$  and  $B$  are  $2 \times 2$  matrices and  $W$  is a two-dimensional Brownian process. The exact solution of this SDE is given by (Arnold, 1974)

In this manuscript, we would calibrate the model to fit with measured data, and compare the effect of using synthetic data compared to measured data for the temperature profile computed by Simstrat. In the long run, we would like to test different wind scenarios, in which case the drift and diffusion terms become time-dependent, that is

$$dZ = A(t)Zdt + B(t)dW \quad (24)$$

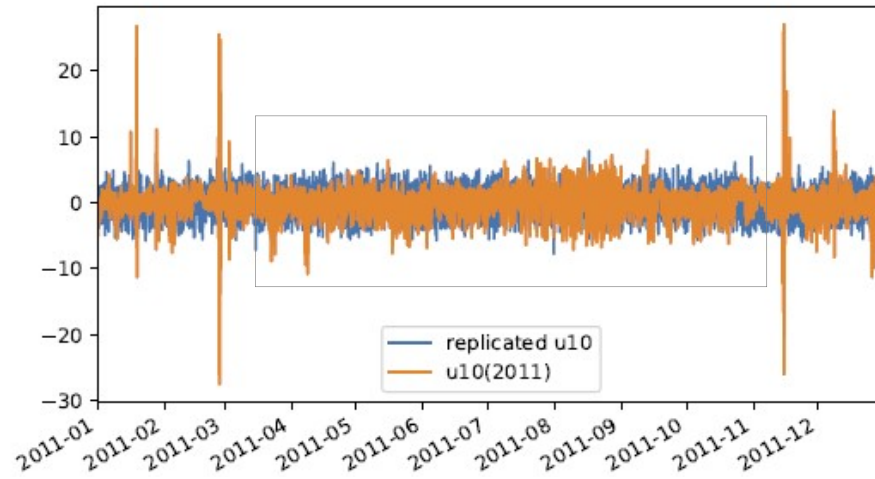
This can be solved with the well known Euler-Maruyama method (Kloeden and Platen, 1992), which for (24)

$$Z_{i+1} = Z_i + hA(t_i)Z_i + B(t_i)\sqrt{h}\xi_i, \quad (25)$$

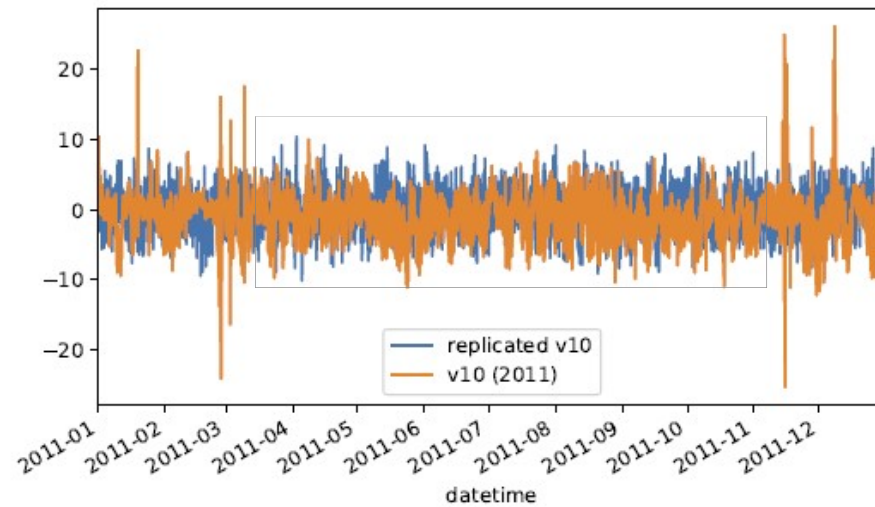
where  $\xi_i \sim \mathcal{N}(0, I_2)$ .

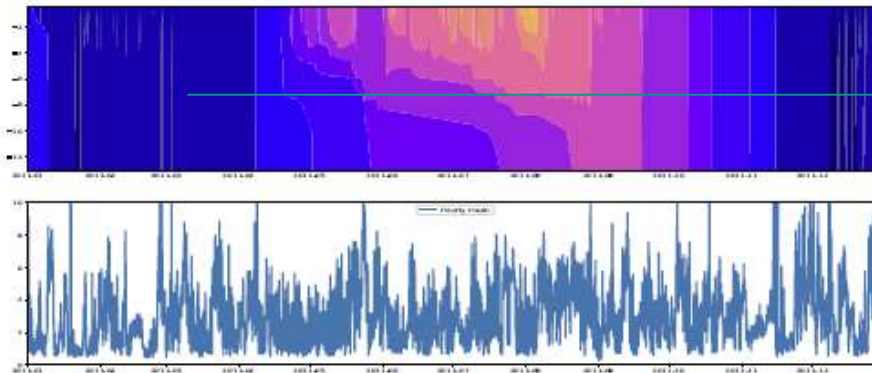


East-West

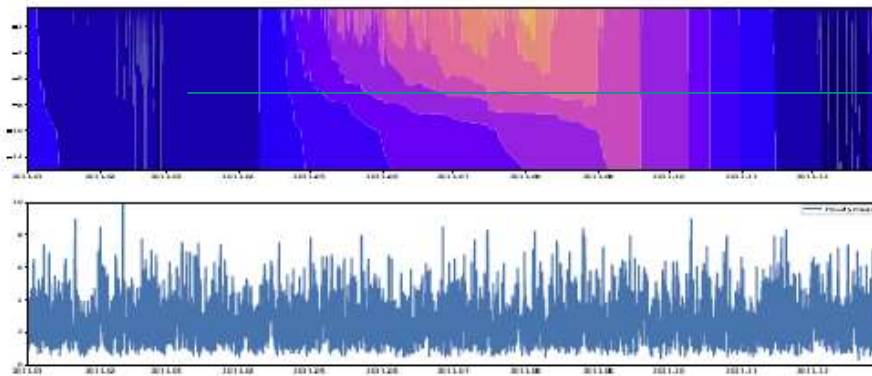


North-South





(a) Temperature profile using measured wind data



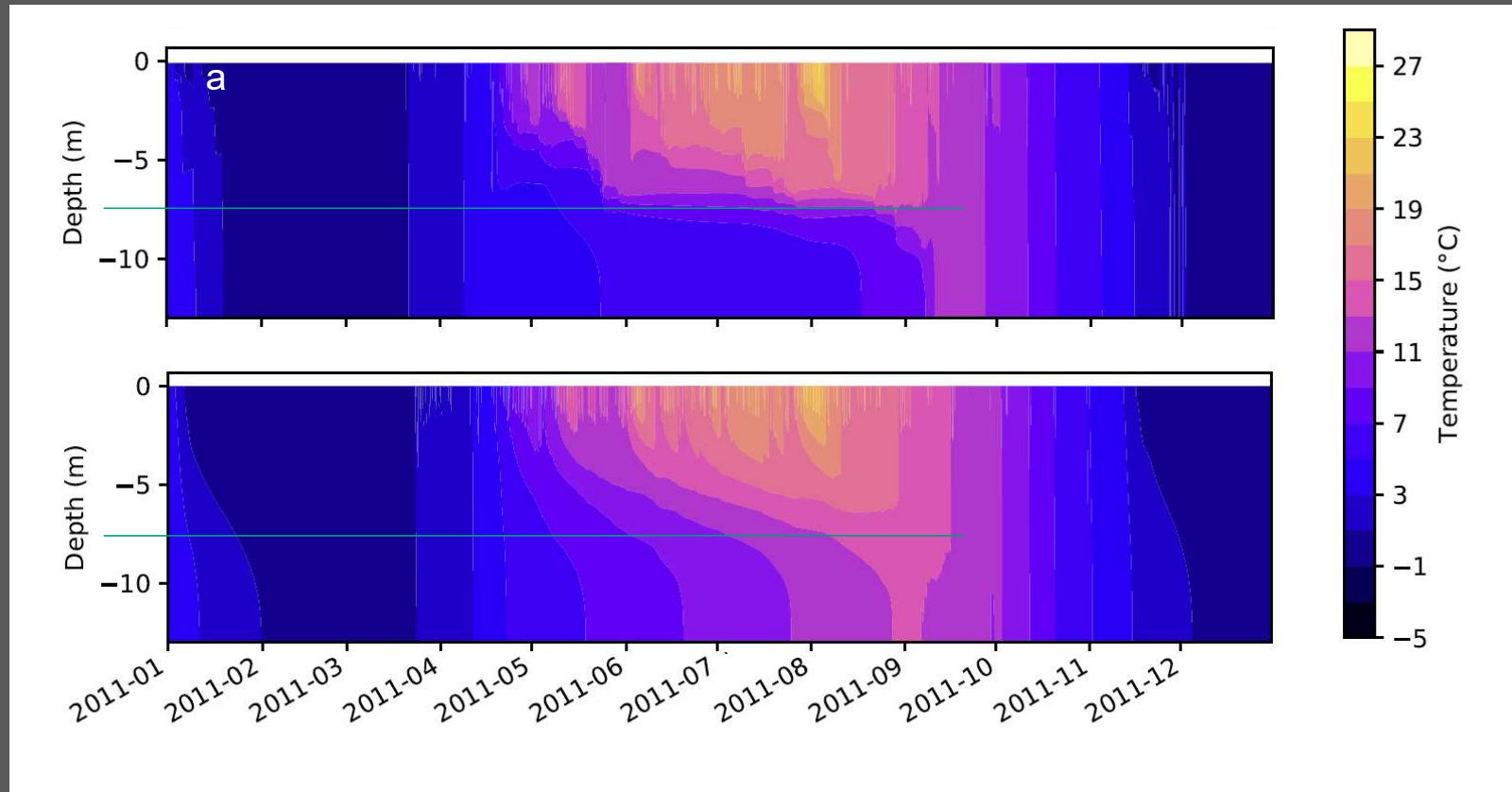
(b) Temperature profile using one simulated wind field

Observed  
wind

Synthetic  
wind



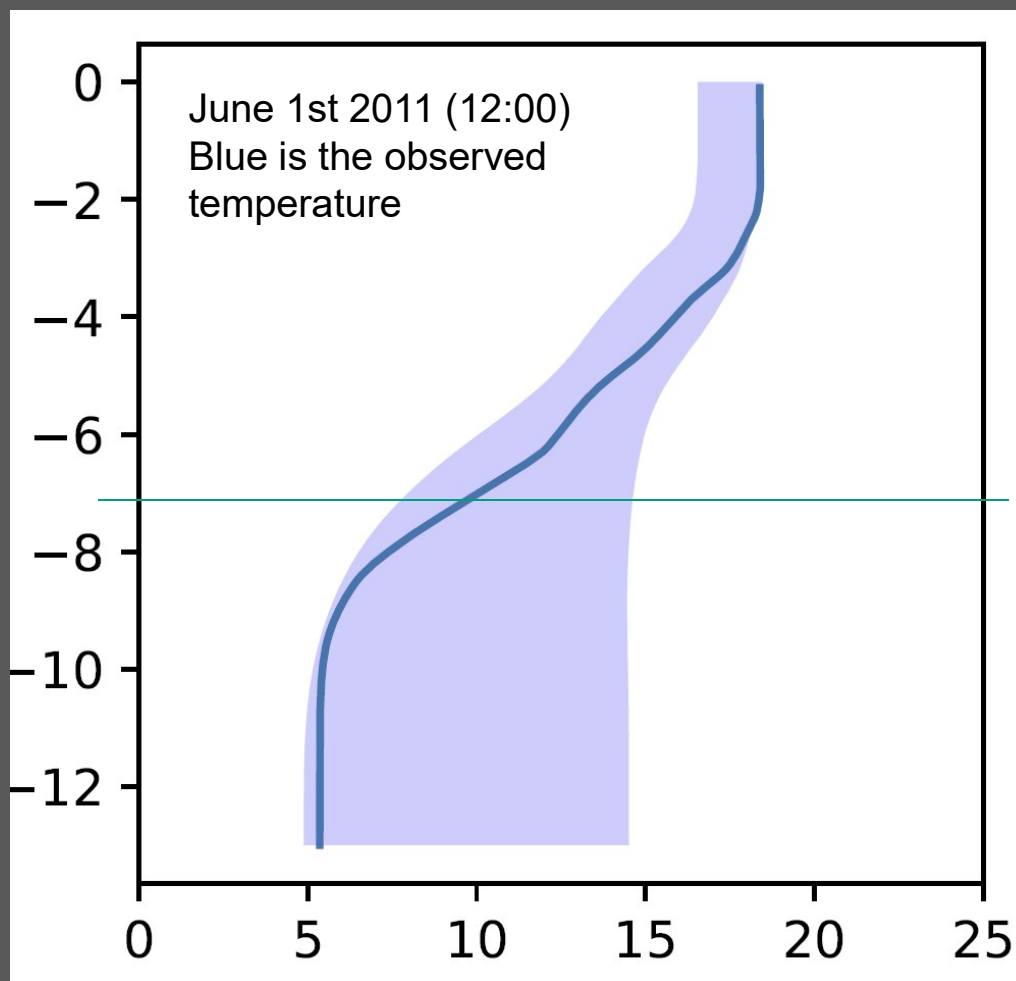
# Temperature simulations





Less variance

More variance



Epilimnion

Metalimnion

Hypolimnion

Probability density function of temperature (T) given depth (z):  $f(T|z)$

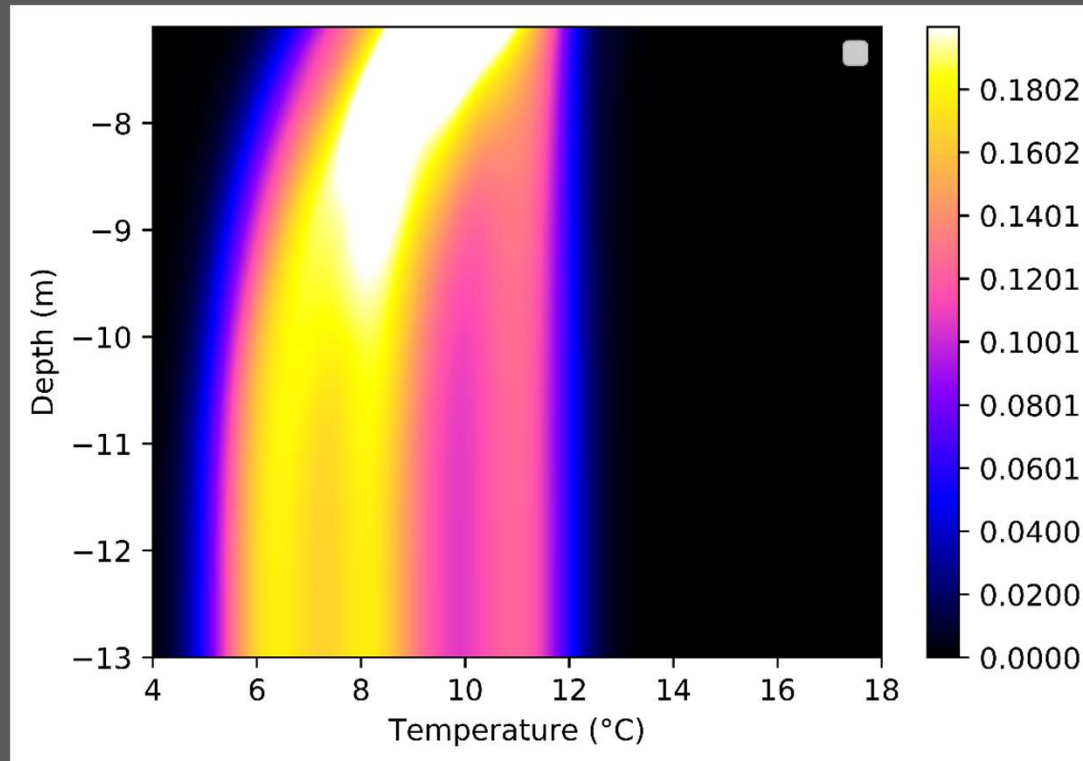


Monivariate pdf ?

Bivariate pdf ??

Multivariate pdf ???

Hypolimnion





## Questions for further work:

- 1) What's the impact of wind to temperature stratification in lakes?
- 2) Do we expect more circulation because of more wind or vice versa?
- 3) Do we need wind observations to mimic temperature in lakes, or can a stochastic wind model do the job?
- 4) Is a simple stochastic model sufficient (mean and covariance), or do we need to include more physics (e.g. coupling of temperature)?
- 5) What time resolution is necessary? Challenge: Mixing parameters depends on variance in wind velocities.



## Key publications:

- Schmid, M.; Ostrovsky, I.; McGinnis, D. F. (2017) **Role of gas ebullition in the methane budget of a deep subtropical lake: What can we learn from process-based modeling?**, *Limnology and Oceanography*, 62(6), 2674-2698, [doi:10.1002/lno.10598](https://doi.org/10.1002/lno.10598).
- Gaudard, A.; Schwefel, R.; Vinnå, L. R.; Schmid, M.; Wüest, A.; Bouffard, D. (2017) **Optimizing the parameterization of deep mixing and internal seiches in one-dimensional hydrodynamic models: a case study with Simstrat v1.3**, *Geoscientific Model Development*, 10(9), 3411-3423, [doi:10.5194/gmd-10-3411-2017](https://doi.org/10.5194/gmd-10-3411-2017).
- Råman Vinnå, L.; Wüest, A.; Bouffard, D. (2017) **Physical effects of thermal pollution in lakes**, *Water Resources Research*, 53(5), 3968-3987, [doi:10.1002/2016WR019686](https://doi.org/10.1002/2016WR019686).
- Goudsmit, G. -H.; Burchard, H.; Peeters, F.; Wüest, A. (2002) **Application of k- $\epsilon$  turbulence models to enclosed basins: the role of internal seiches**, *Journal of Geophysical Research C: Oceans*, 107(C12), 3230 (13 pp.), [doi:10.1029/2001JC000954](https://doi.org/10.1029/2001JC000954).
- Fink, G.; Schmid, M.; Wüest, A. (2014) **Large lakes as sources and sinks of anthropogenic heat: capacities and limits**, *Water Resources Research*, 50(9), 7285-7301, [doi:10.1002/2014WR015509](https://doi.org/10.1002/2014WR015509).